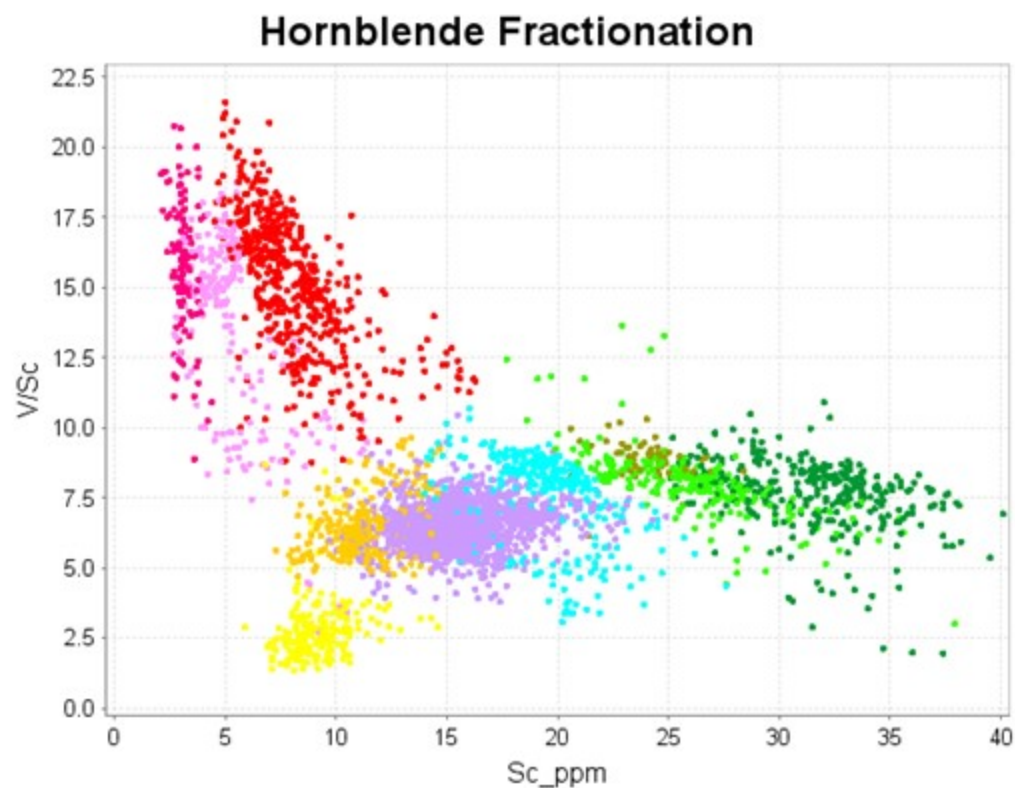


Understanding magmatic fractionation processes using multielement geochemistry

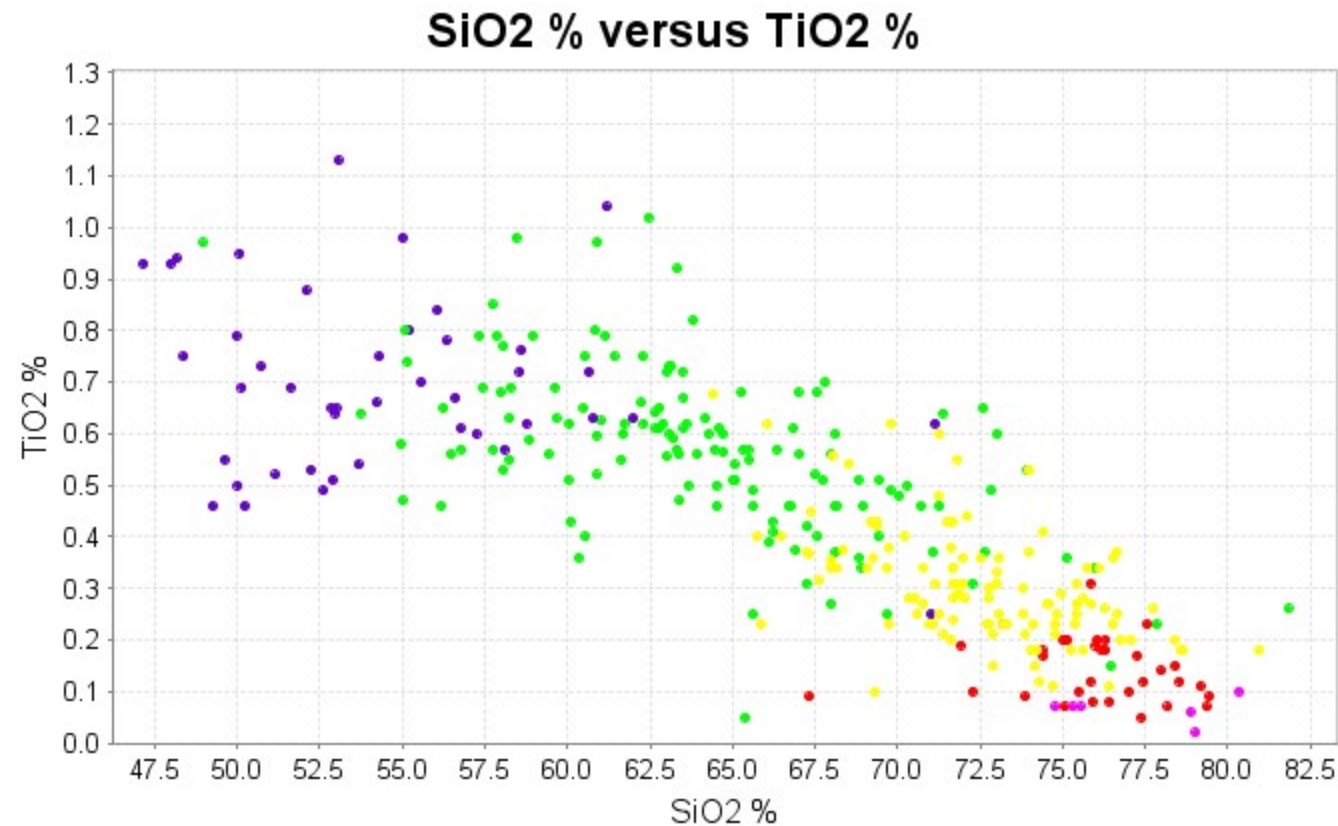


Summary

- Fractional crystallization; segregation of crystals from a melt in a cooling magma.
- Concentrates volatiles in the melt.
- Volatile saturation produces a hydrothermal system and hopefully an orebody.
- This presentation
 - low pressure, mid-crustal melts
 - compare fractionation from high-pressure hydrous melting (eg porphyry Cu)
 - Diversity of Archean granites

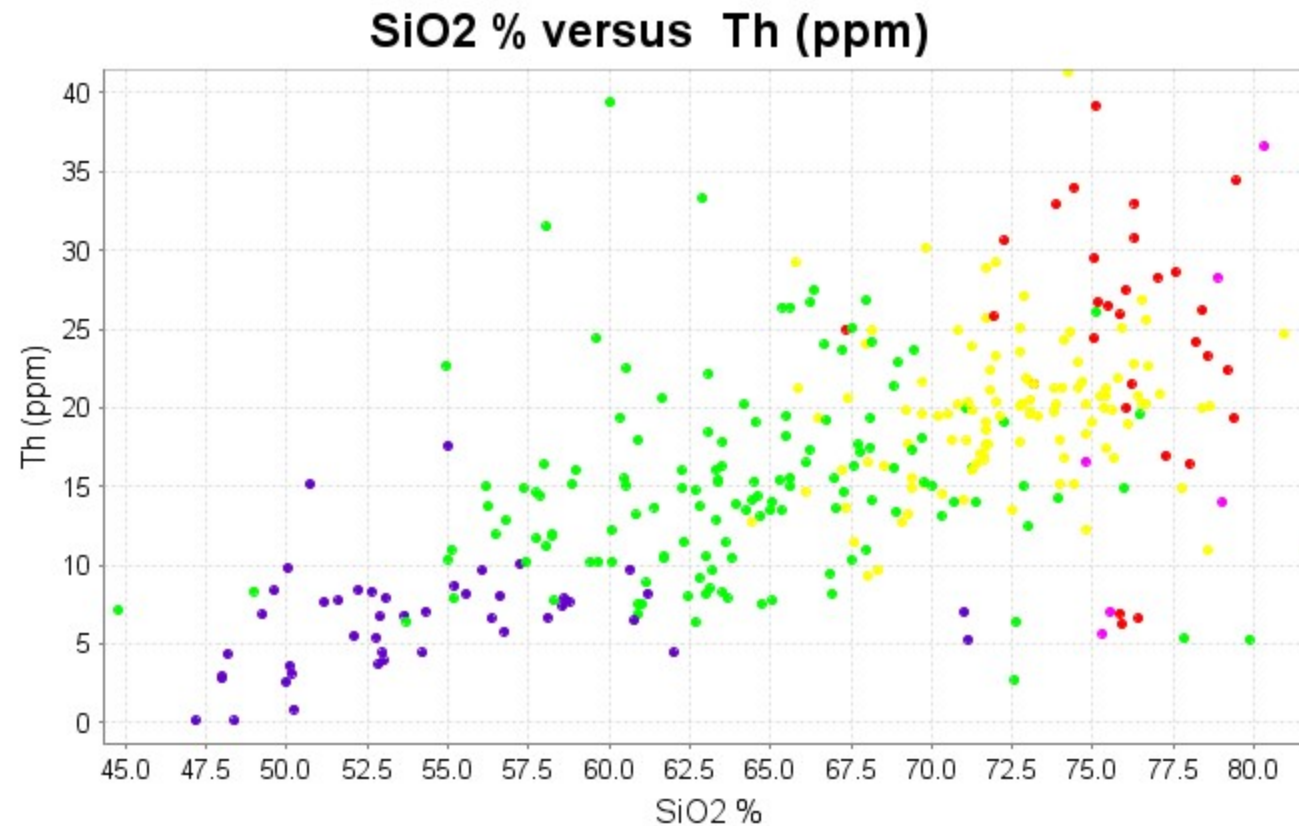
Compatible elements have decreasing abundance with increasing SiO₂

Example: Ti is a *compatible* element.



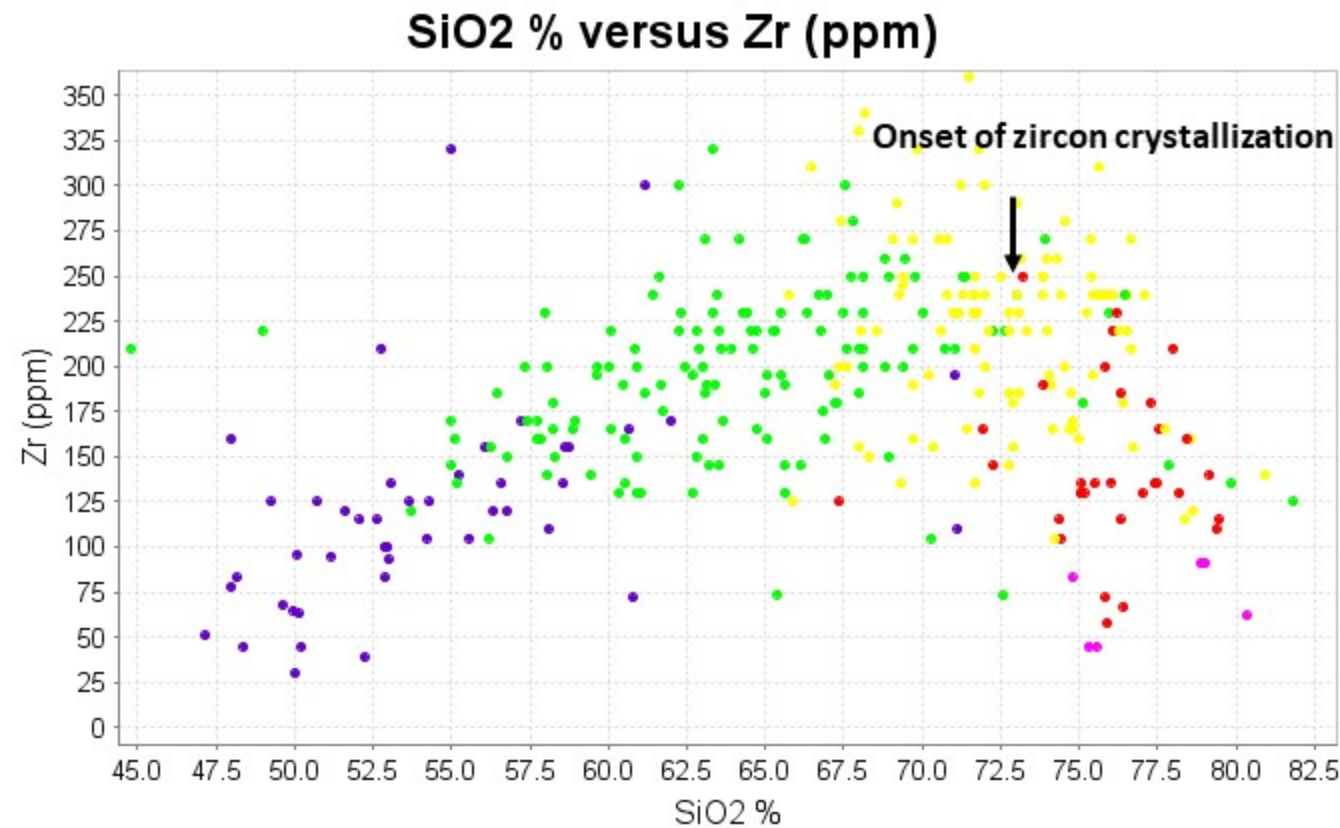
Incompatible elements have increasing abundance with increasing SiO₂

Example: Th is an *incompatible* element.



Accessory mineral phases control HFSE distribution patterns

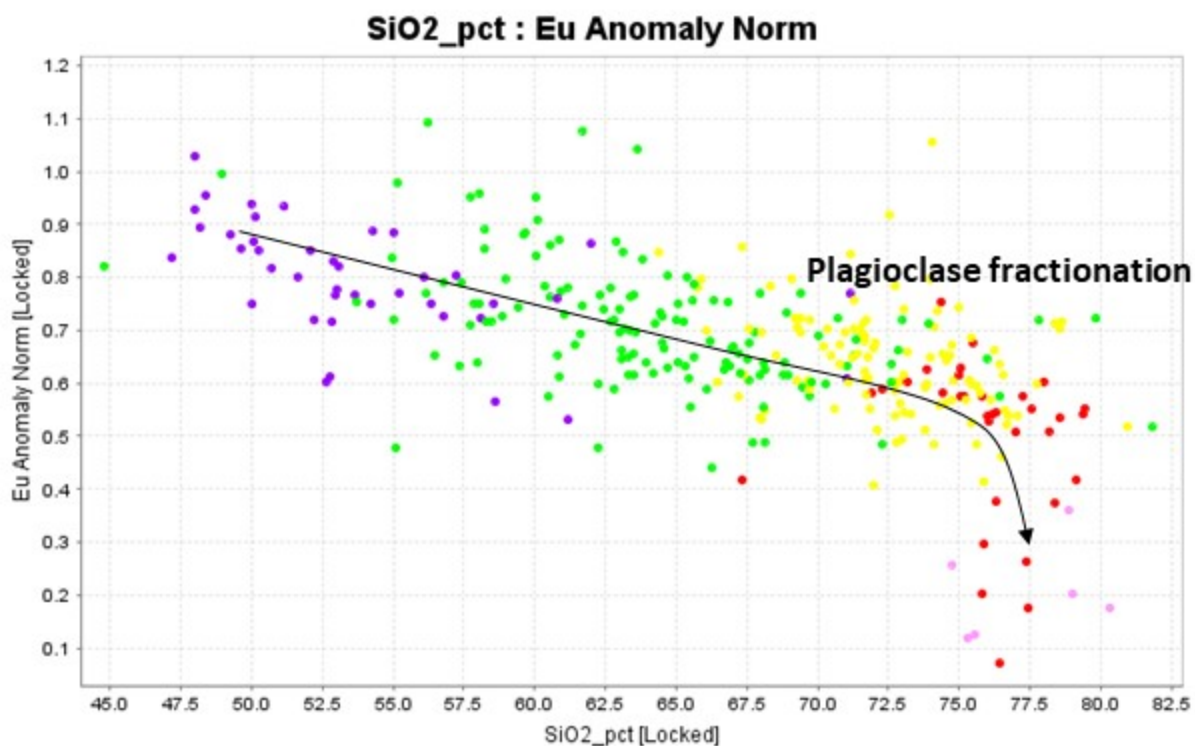
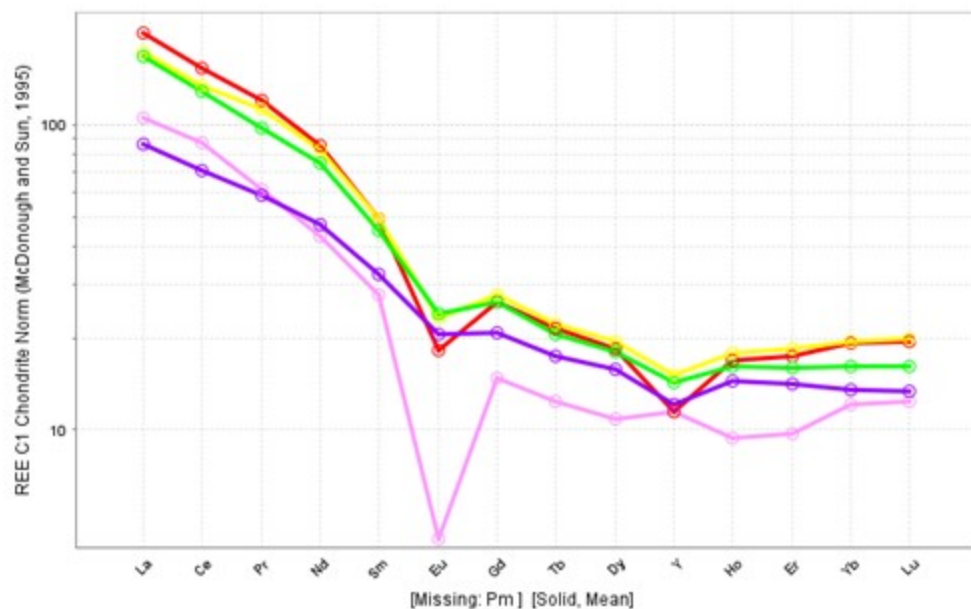
Example; zircon crystallization begins at around 800°C; Zr changes from incompatible to compatible behavior at around 70% SiO₂



Fractionation in low pressure (crustal) hydrous melts

Plagioclase Fractionation

REE's valency 3⁺. Eu can be 2⁺. In reduced to moderately oxidized melts, Eu²⁺ substitutes for Ca²⁺. Plagioclase fractionation removes Eu from a melt relative to other REE's.



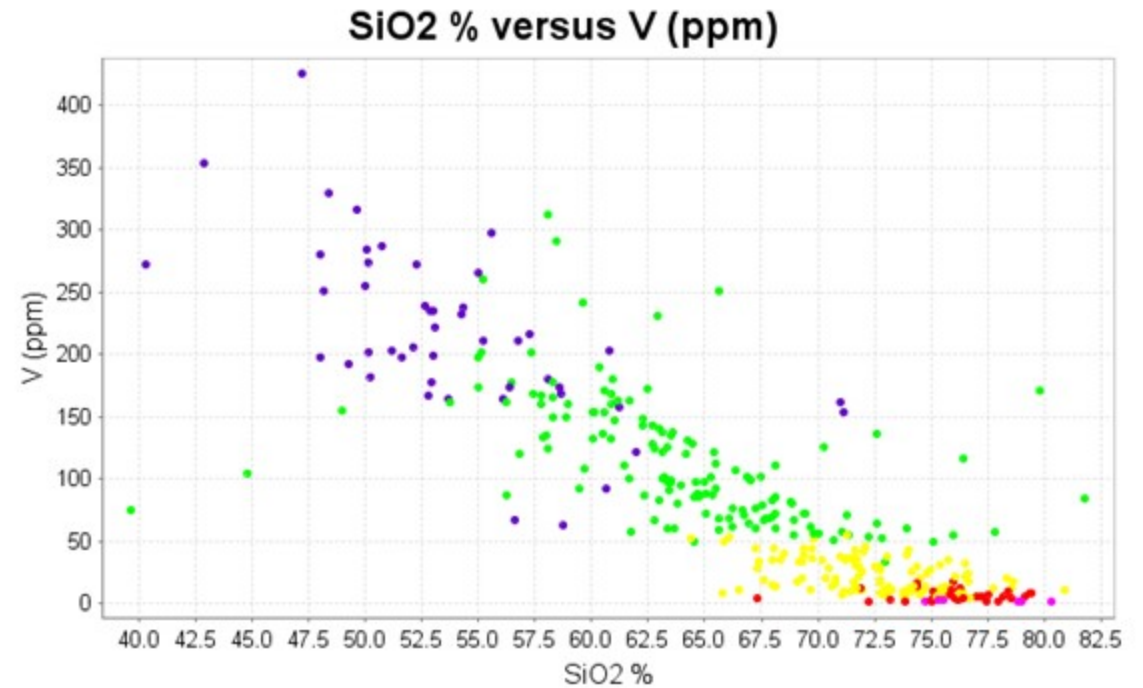
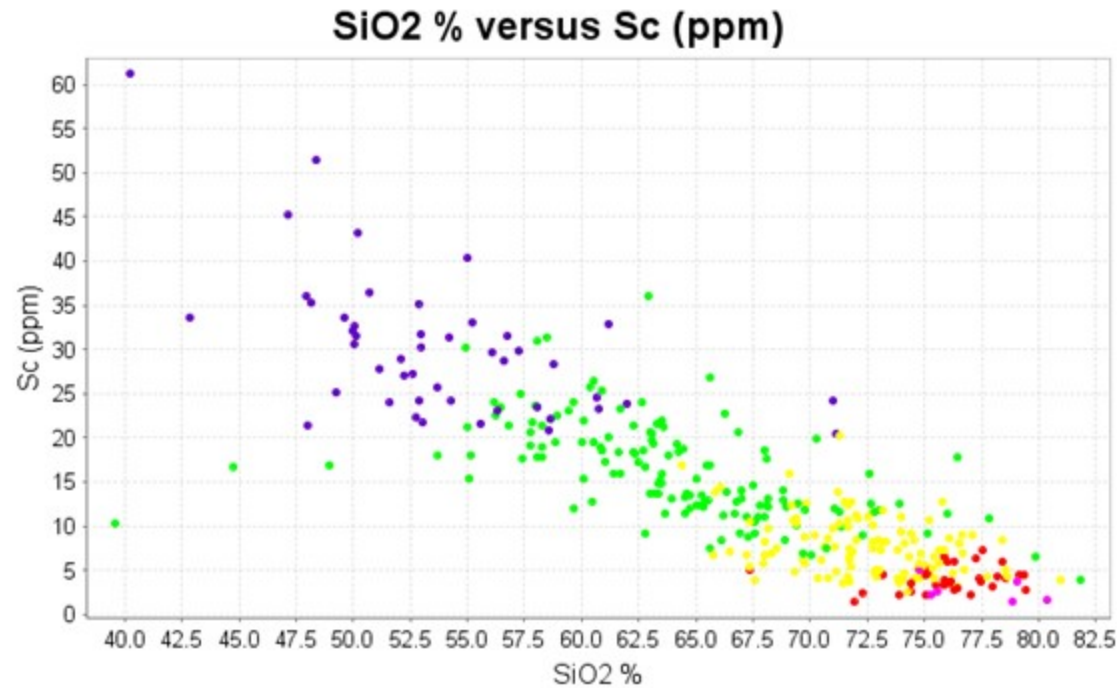
Fractionation in low pressure (crustal) hydrous melts

Trace Element Pairs

Sc and V^{3+} are both compatible elements with similar partition coefficients. Sc and V^{3+} substitute for Fe in silicate minerals (behaves like the silicate-hosted component of Fe).

V^{4+} behaves more like Ti; it is less compatible in silicates but more compatible in oxides.

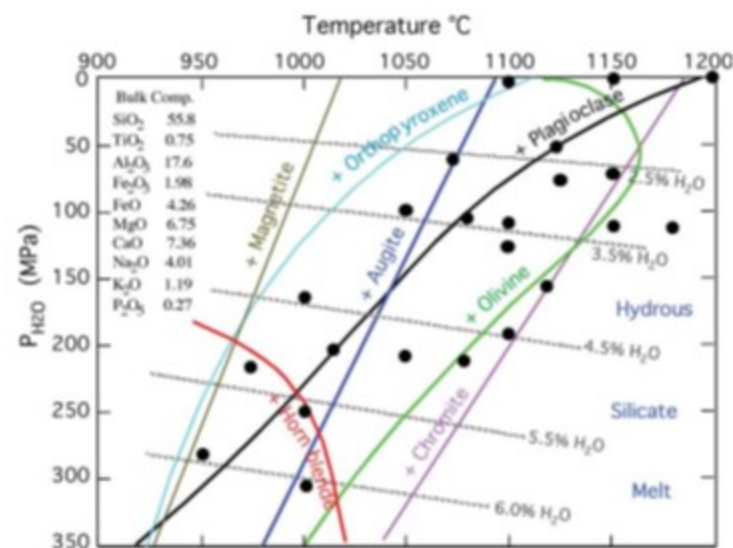
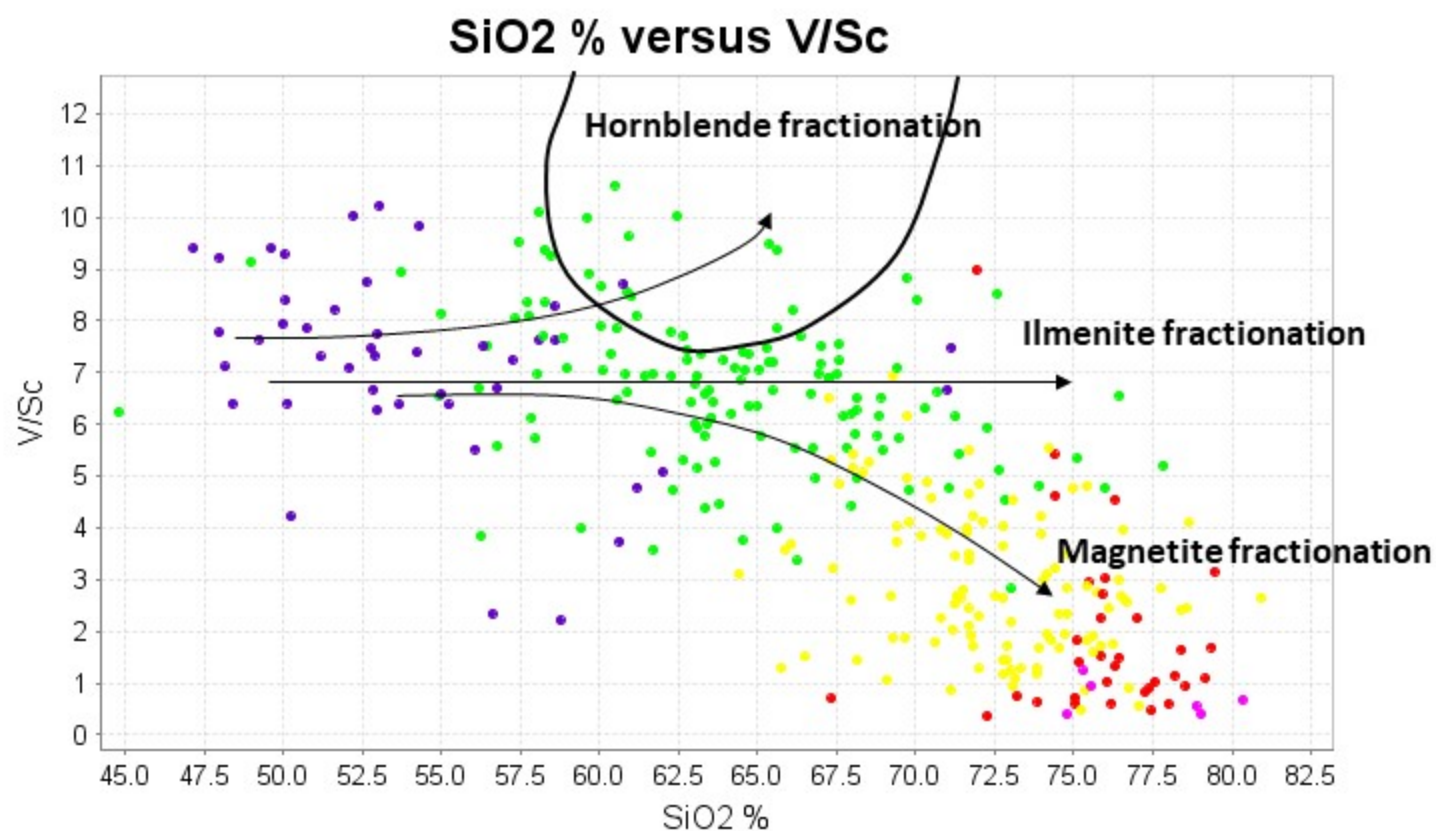
Sc and V have the same trajectories in reduced magma, but different trajectories in oxidized melts.



Fractionation in low pressure (crustal) hydrous melts

Magnetite versus Hornblende fractionation

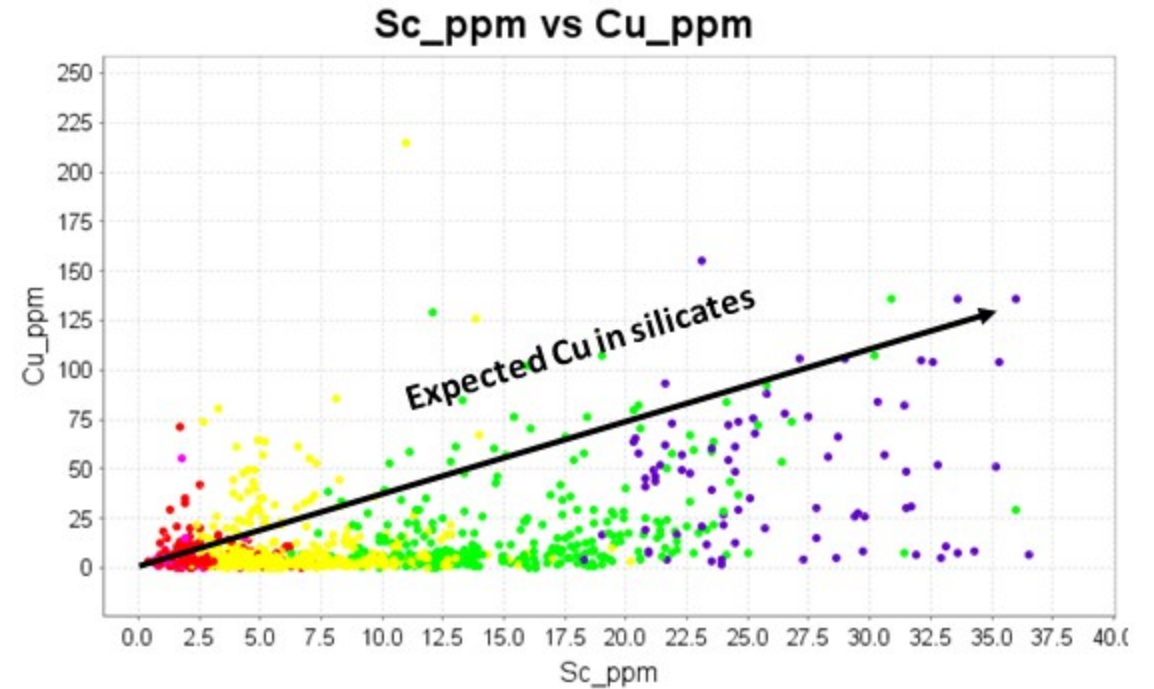
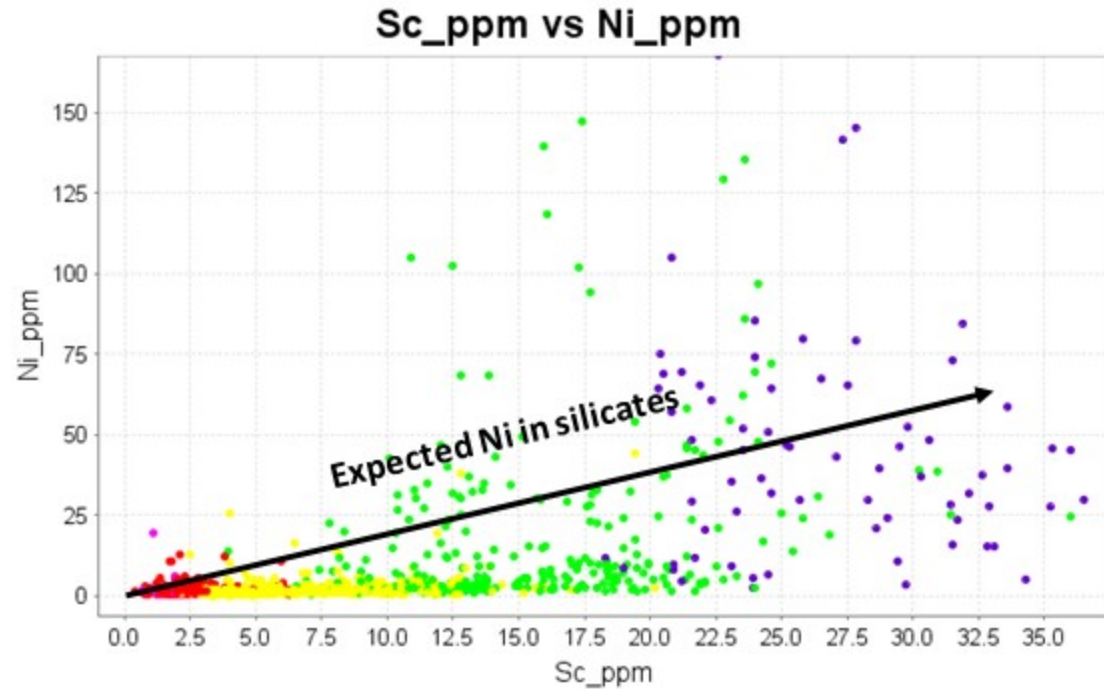
In an oxidized melt (ie some V⁴⁺), if magnetite crystallizes before hornblende, then V is removed from the melt faster than Sc, and the V/Sc ratio declines. In an oxidized melt where V⁴⁺ is somewhat incompatible in silicates, crystallization of hornblende before magnetite causes the V/Sc ratio to increase. The pattern shown in this figure is typical of low pressure, magnetite-bearing crustal melts.



Fractionation in low pressure (crustal) hydrous melts

Sulfide Saturation

Magmas that fractionate magnetite are invariably sulfide saturated. This is evident from depleted Ni and Cu relative to Sc.

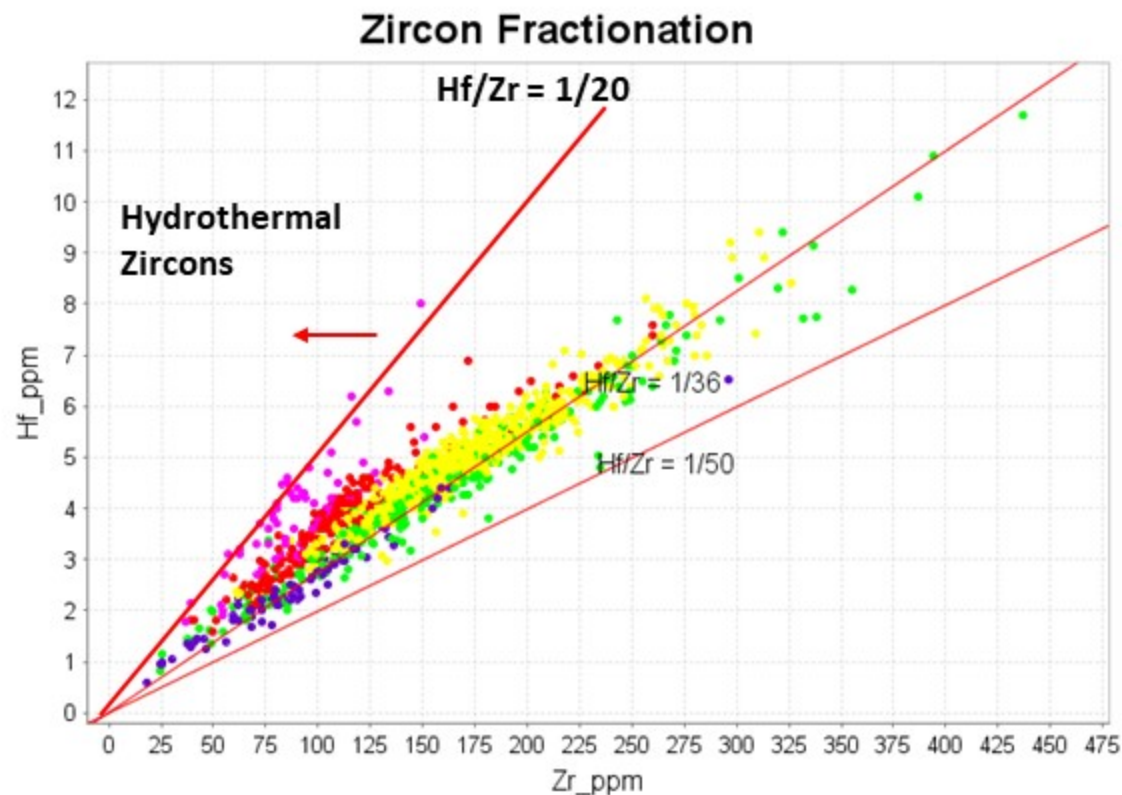
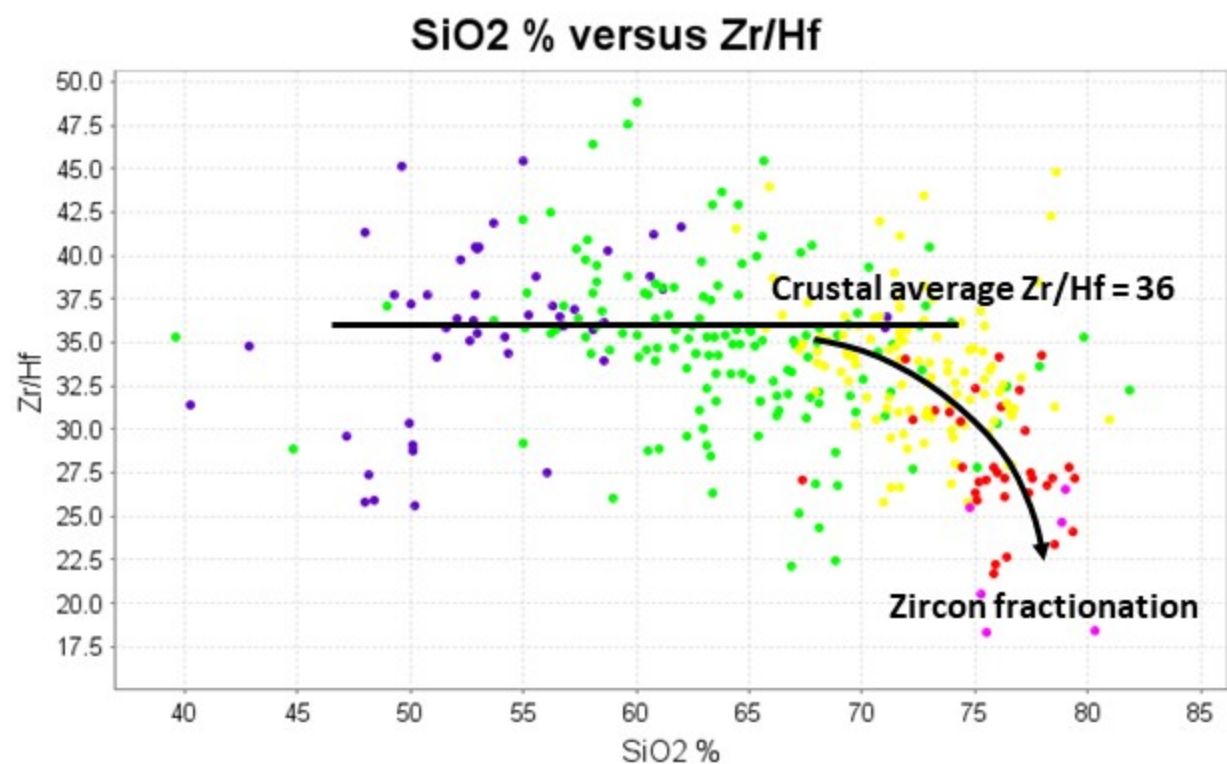


Fractionation in low pressure (crustal) hydrous melts

Accessory mineral phases control HFSE distribution patterns

Example; Hf always follows zirconium. K_D for Hf in Zircons is <1 ,

Therefore zircon fractionation causes Zr/Hf ratio to decrease. (at $> 70\%$ SiO_2)



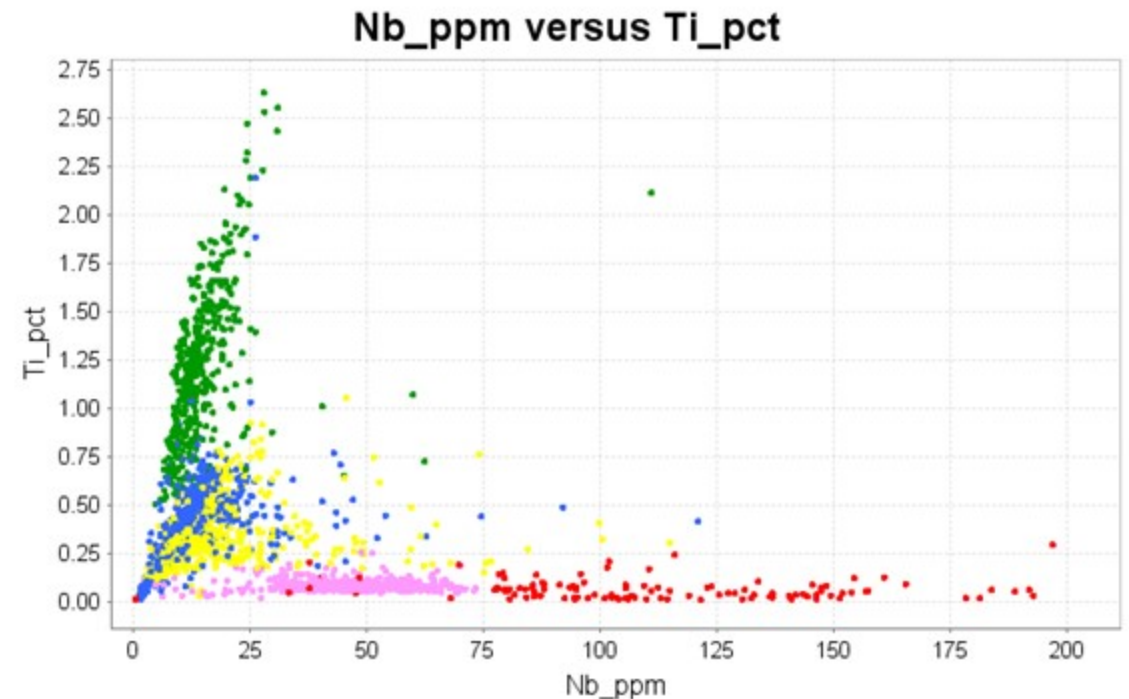
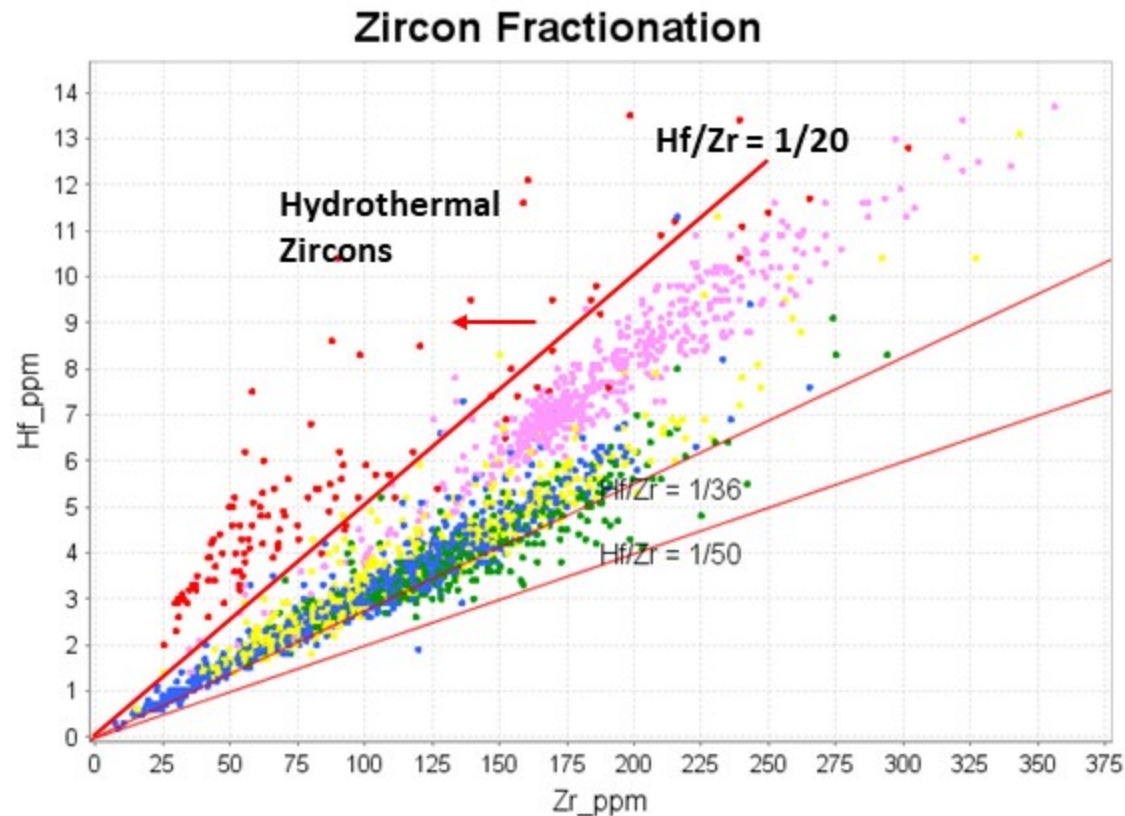
Fractionation in low pressure (crustal) hydrous melts

Zircon Zr/Hf ratio; example from a major Sn granite deposit

Most rocks have $\text{Zr}/\text{Hf} = 36$

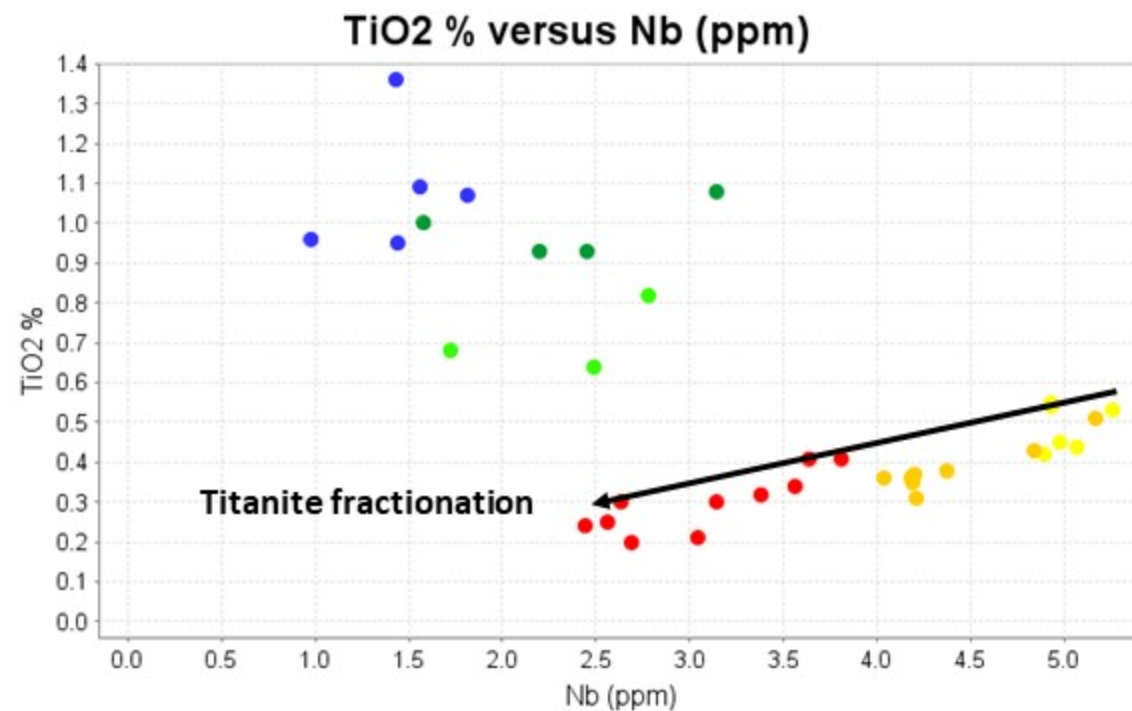
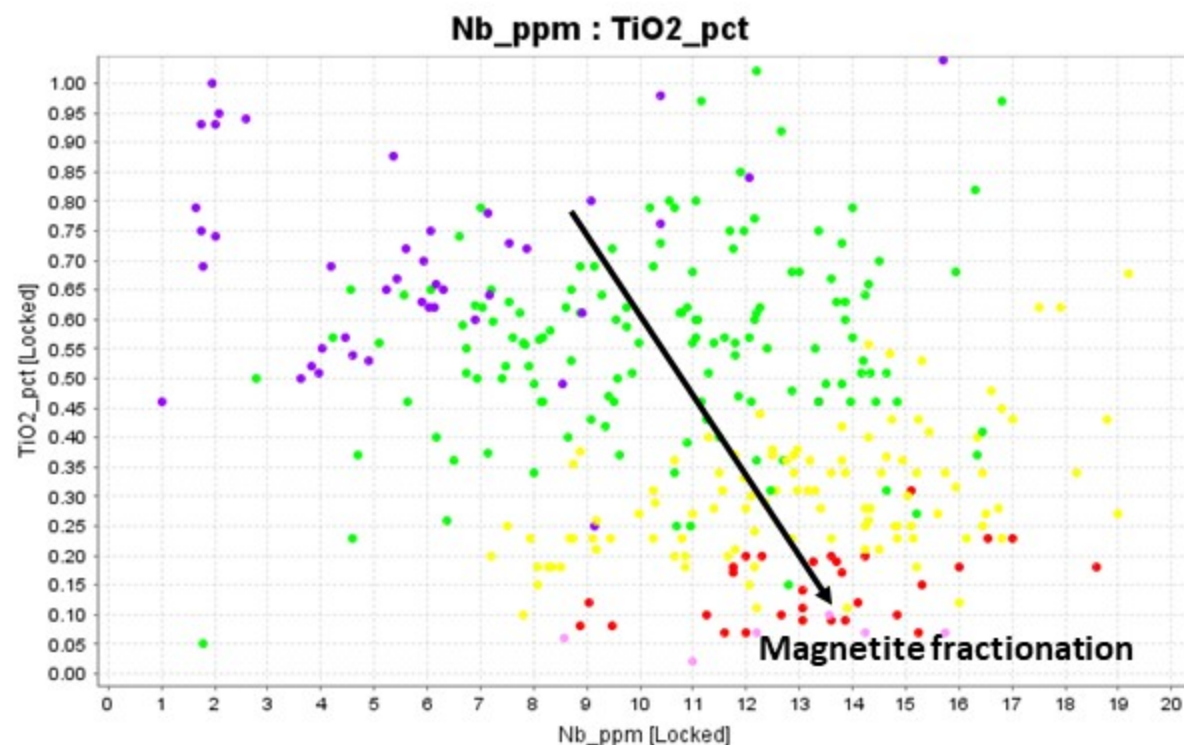
Zircon fractionation at $\text{WR SiO}_2 > 70\%$ Zr/Hf goes from 36 to 20

Hydrothermal zircons, eg in pegmatites and aplites $\text{Zr}/\text{Hf} < 20$



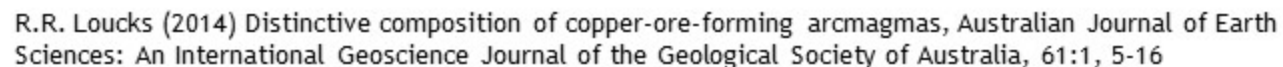
Fractionation in low pressure (crustal) hydrous melts

Ti versus Nb tells you a lot about the opaque oxide minerals. In crustal melts, Nb increases as Ti decreases.



Hornblende fractionation leads to increasing V/Sc with increasing SiO₂.

This is a diagnostic characteristic of magmas associated with porphyry Cu systems.



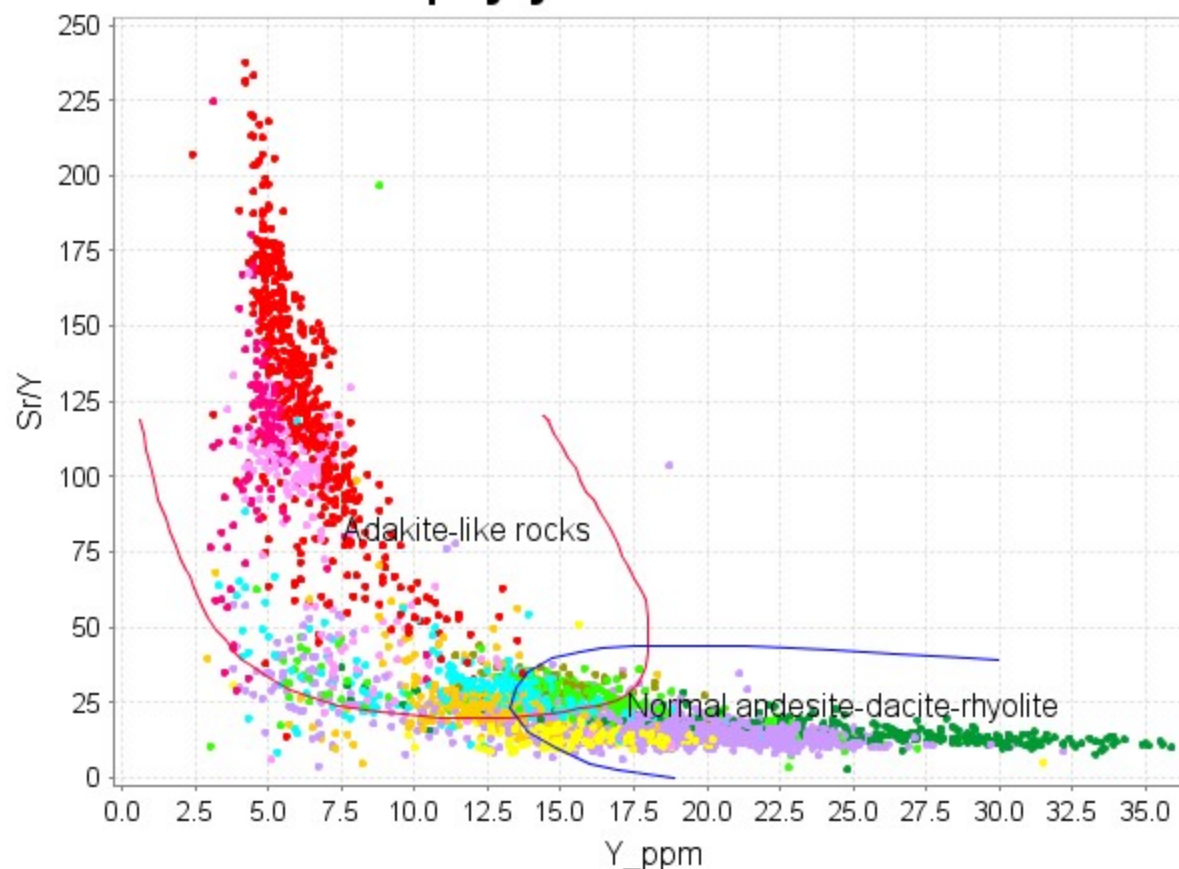
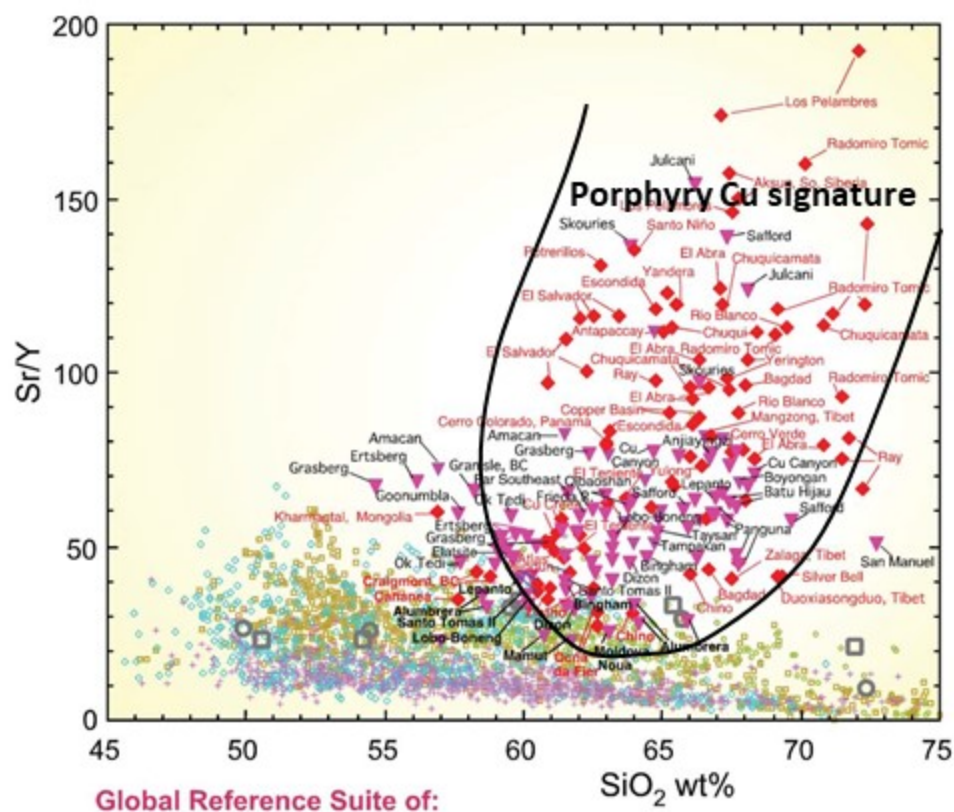
Fractionation in high pressure hydrous melts (porphyry Cu systems)

Strontium versus Yttrium

Garnet-stable source region retains Y => low Y magma

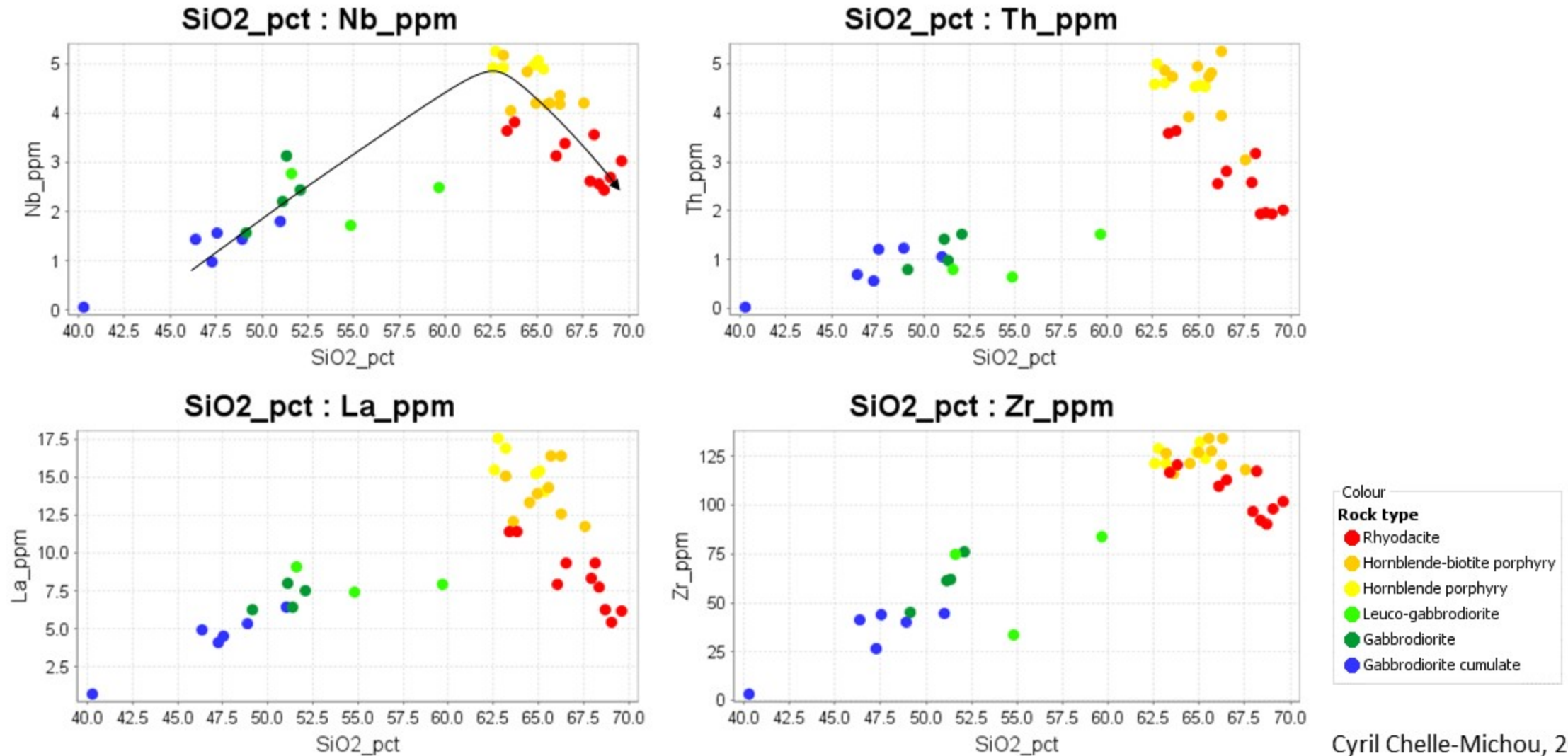
High water content lowers the plagioclase liquidus; delayed onset of plag crystallization = high Sr magma

Porphyry Cu Sr/Y vs Y

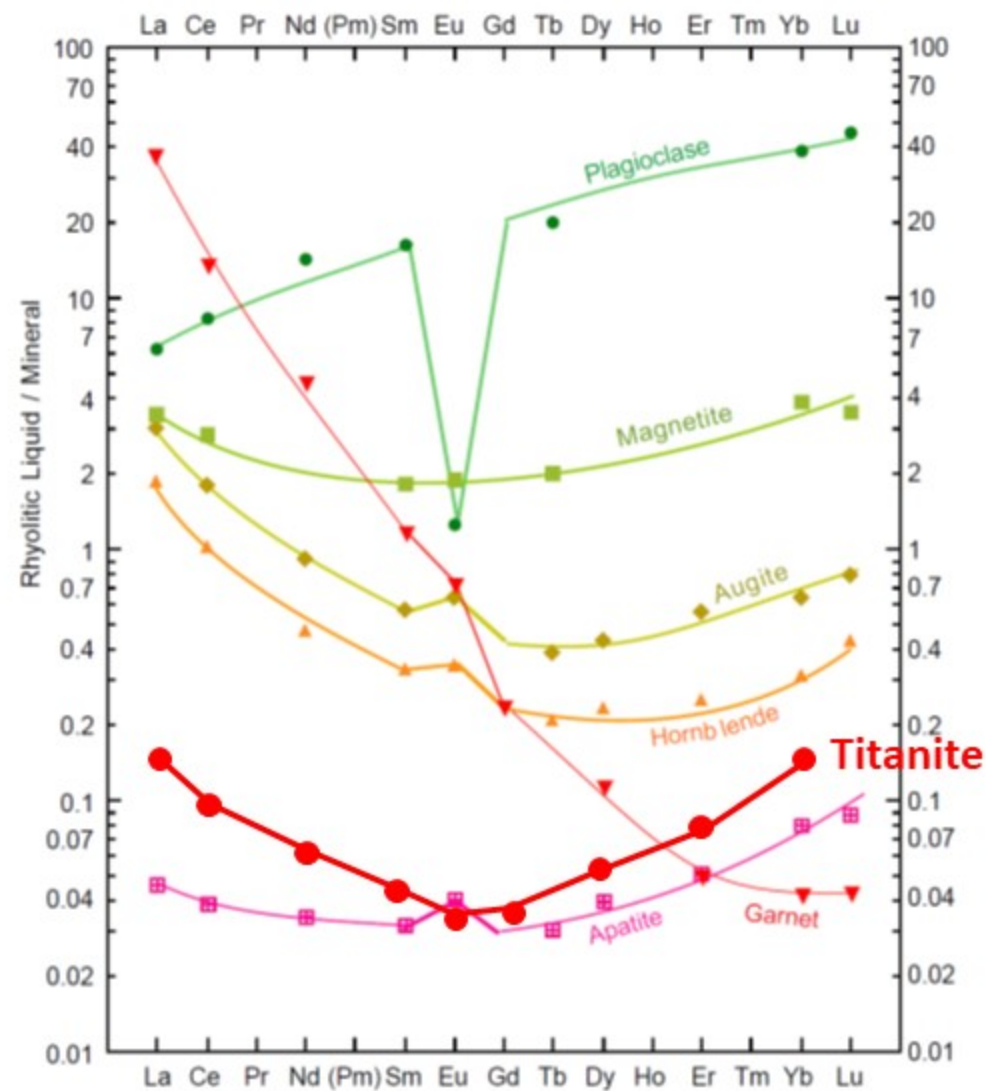
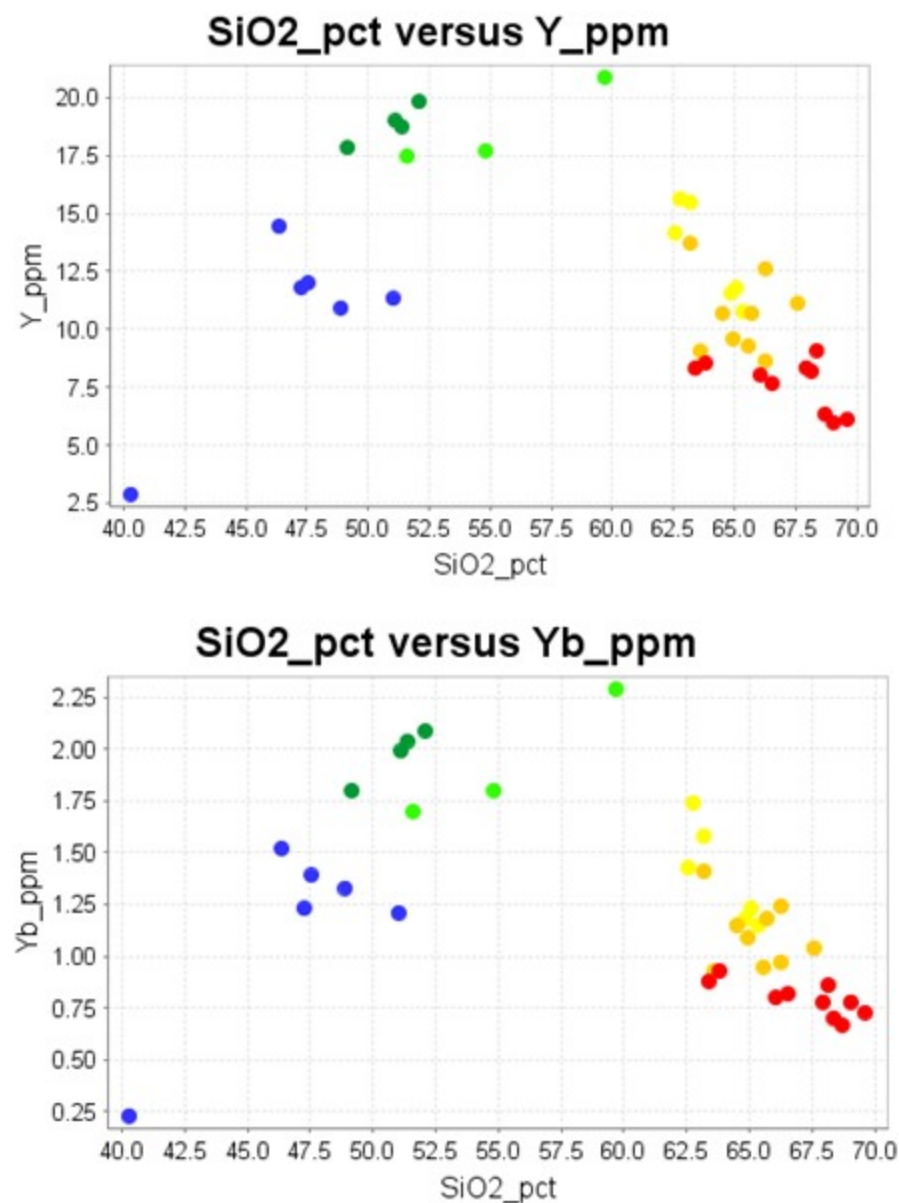


The role of titanite (CaTiSiO_5) in HFSE patterns in porphyry Cu systems

The onset of titanite crystallization in porphyry Cu magmas causes “compatible” behaviour of Nb, Th, REE’s, Zr, Y and U above ~62% SiO_2



Titanite fractionation MUST have a significant impact on whole-rock Y (equivalent to MHREE)



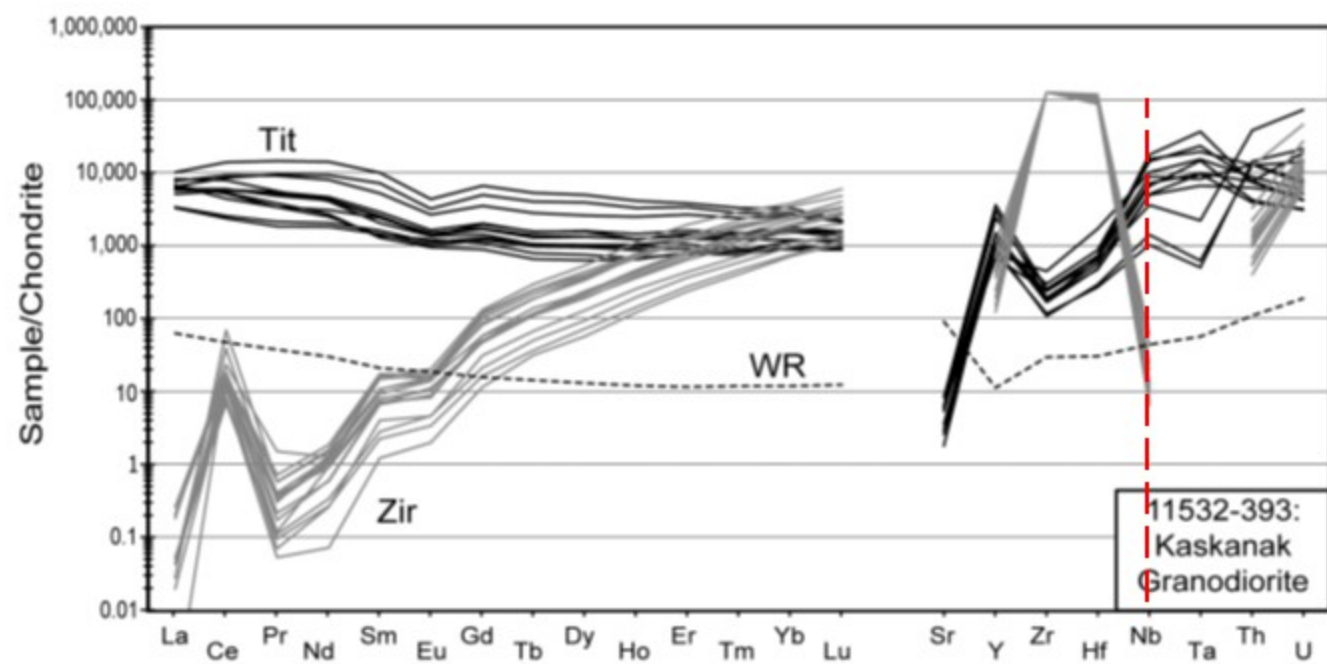
Tiepolo, M., Oberti, R. and Vannucci, R., 2002. Trace-element incorporation in titanite: constraints from experimentally determined solid/liquid partition coefficients. *Chemical Geology*, 191(1-3), pp.105-119.

R.R. Loucks (2014) Distinctive composition of copper-ore-forming arc magmas, *Australian Journal of Earth Sciences: An International Geoscience Journal of the Geological Society of Australia*, 61:1, 5-16

The role of titanite (CaTiSiO_5) in HFSE patterns in porphyry Cu systems

Representative REE and HFSE contents of titanite, zircon, and wholerock normalized to chondrite. Chondrite values from McDonough and Sun (1995)

Fractionation of titanite accounts for the compatibility of Nb, Ta, REE's, Th, Zr, Hf

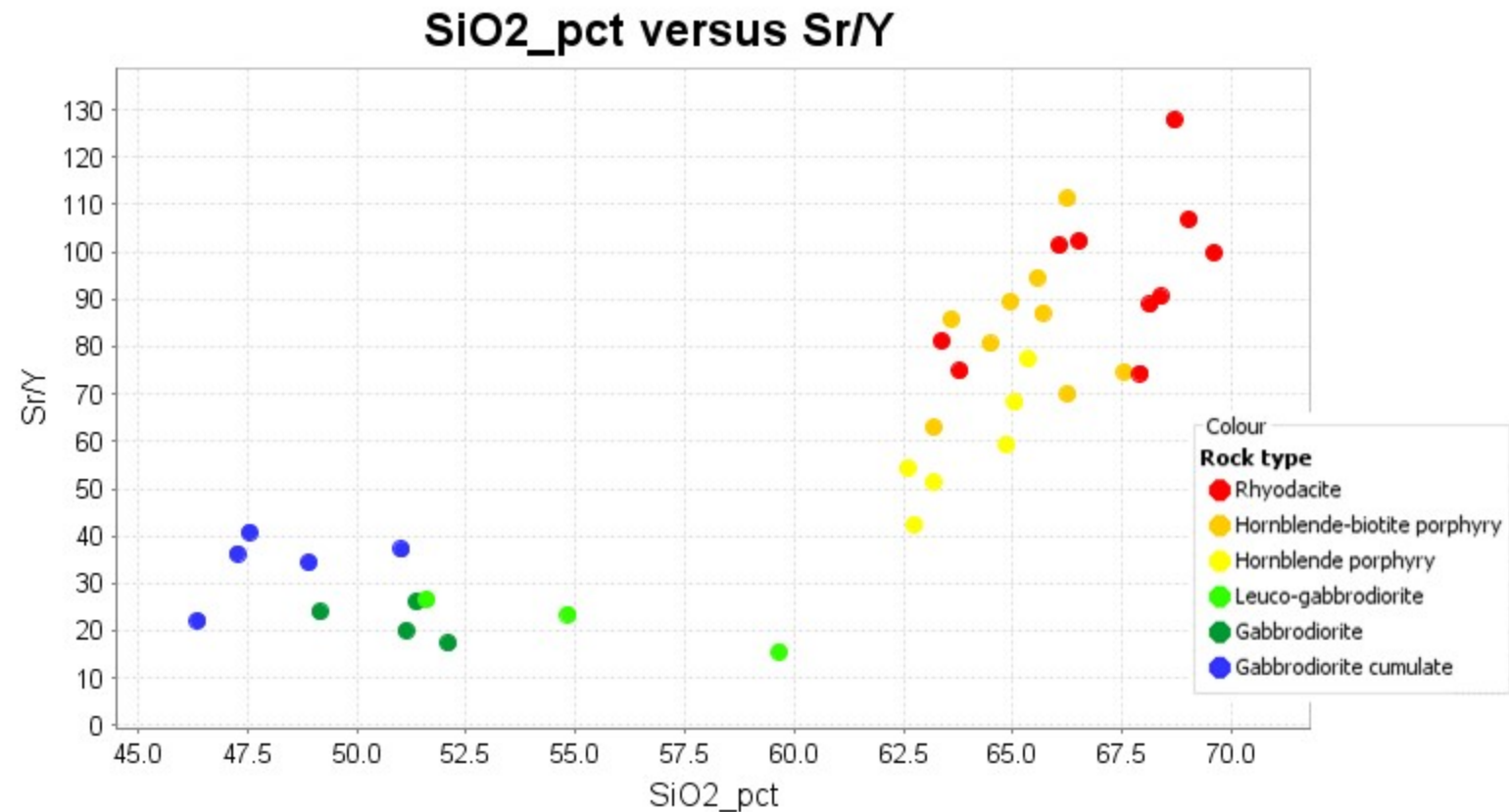


Olson, N.H., Dilles, J.H., Kent, A.J. and Lang, J.R., 2017. Geochemistry of the Cretaceous Kaskanak batholith and genesis of the Pebble porphyry Cu-Au-Mo deposit, southwest Alaska. *American Mineralogist: Journal of Earth and Planetary Materials*, 102(8), pp.1597-1621.

The role of titanite (CaTiSiO_5) in HFSE patterns in porphyry Cu systems

The increase in Sr/Y as porphyry Cu magmas evolve is a result of titanite fractionation rather than the effect of Y retention in garnet at the source region.

The trend of increasing Sr/Y in porphyry Cu magmas is commonly interpreted to result from increasing depth (pressure) at the source of melting.



The role of titanite (CaTiSiO_5) in HFSE patterns in porphyry Cu systems

Sr/Y patterns in different porphyry Cu environments all look the same!!

Oceanic Island Arc porphyries, (thin crust)

Andean porphyries

SW USA, (thick crust)

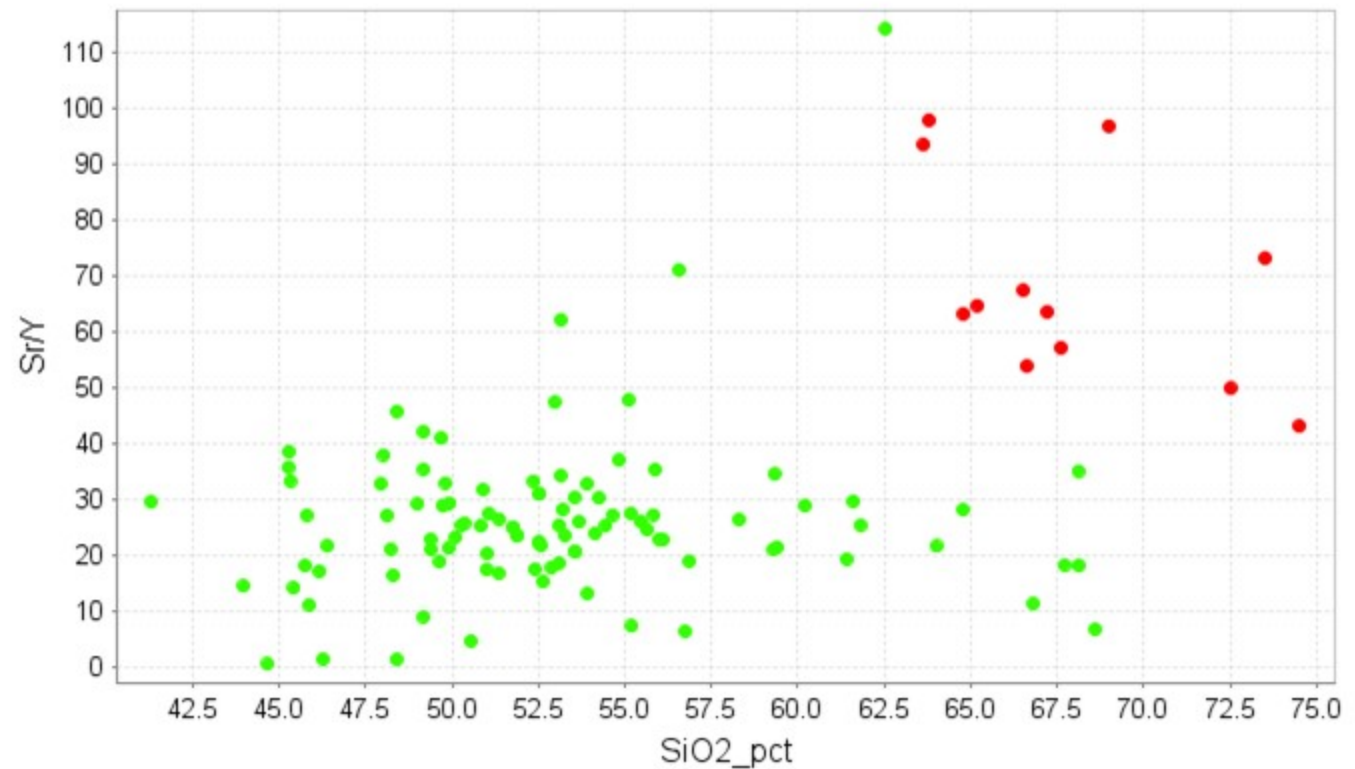
Porphyry Mo (very thick crust)

This example is from the Indonesian Arc

Red = hornblende diorite (porphyry Cu)

Green = basaltic andesite country rock

Indonesian Porphyry Cu; SiO₂ vs Sr/Y

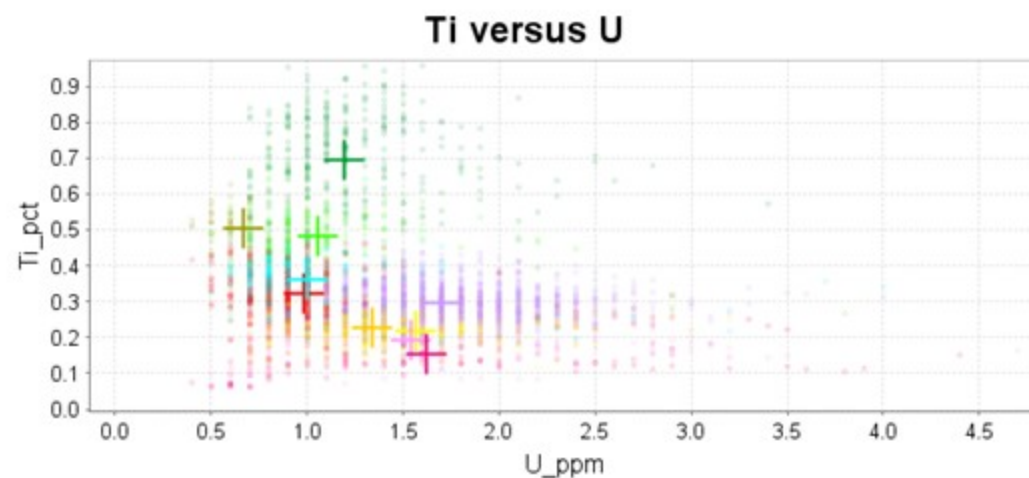
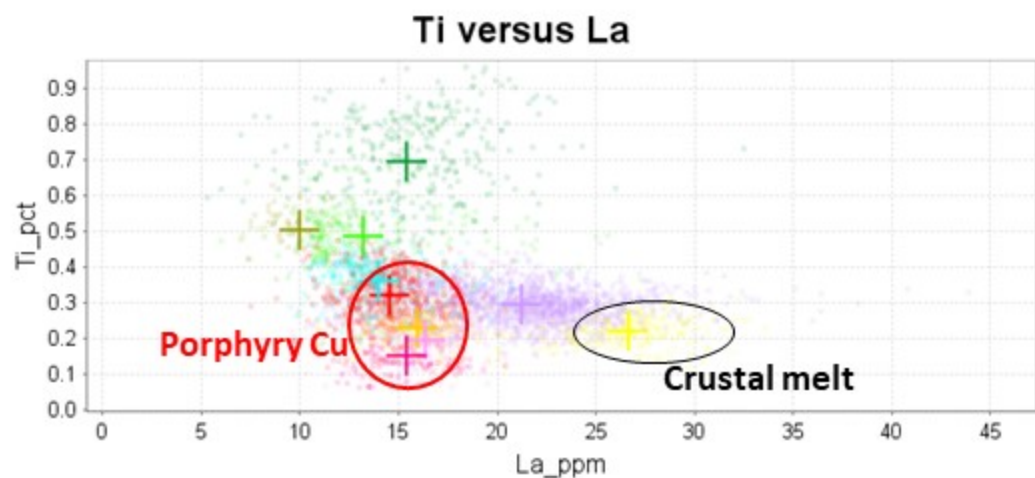
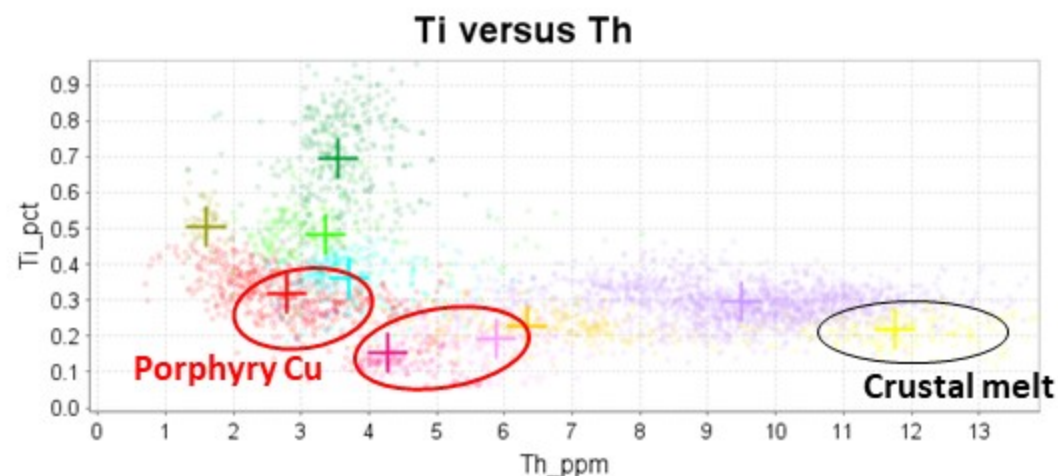
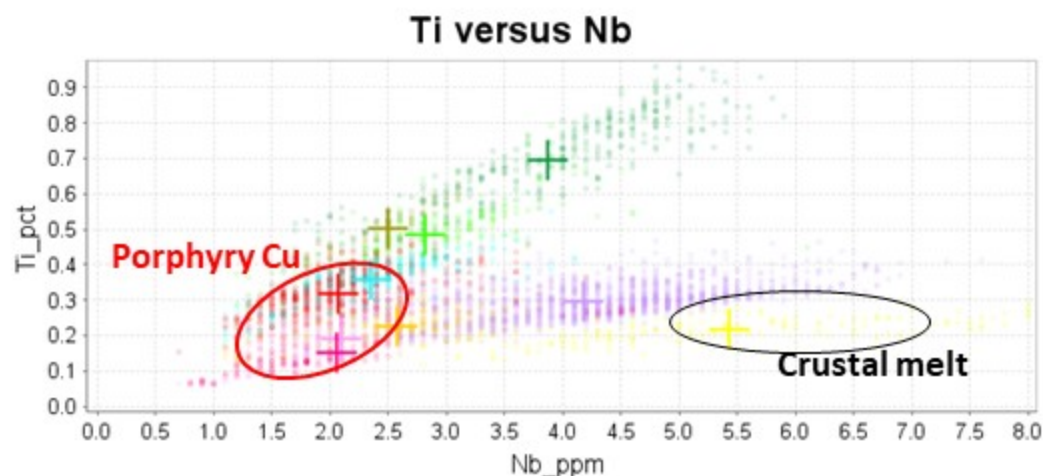


Fractionation in high pressure hydrous melts (porphyry Cu systems)

High Ca magma favours titanite fractionation.

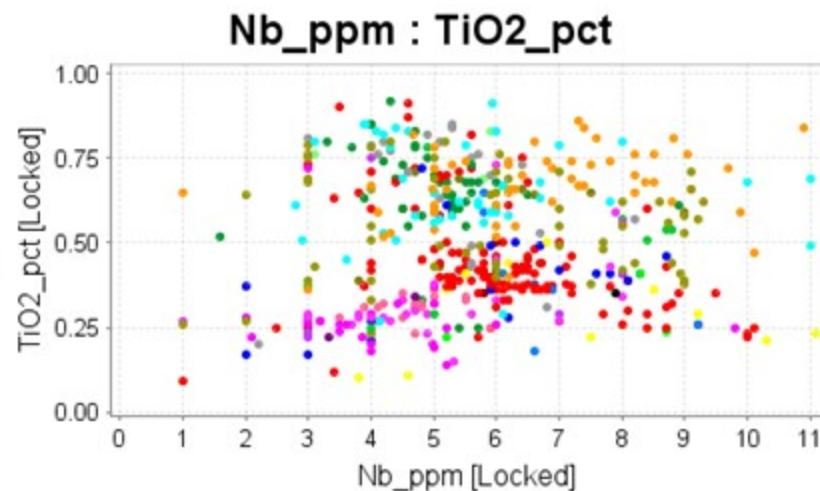
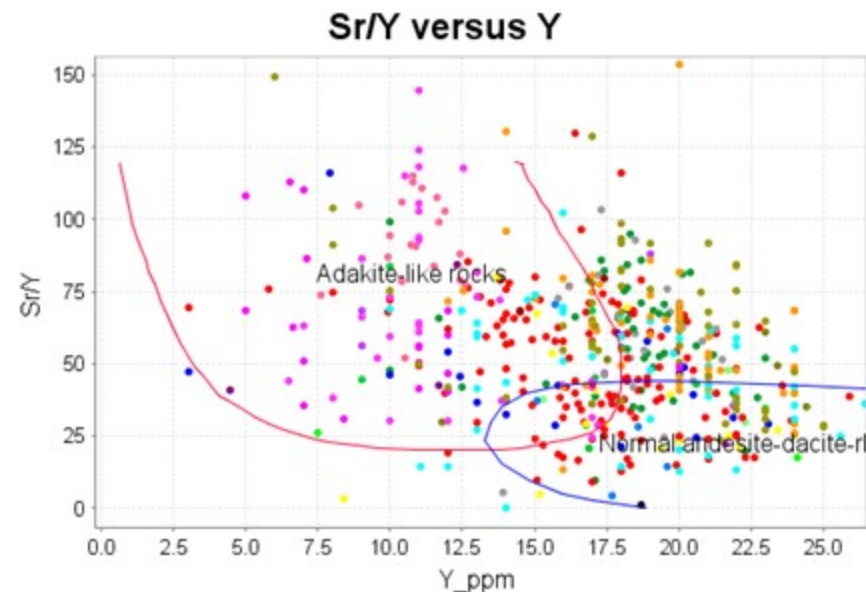
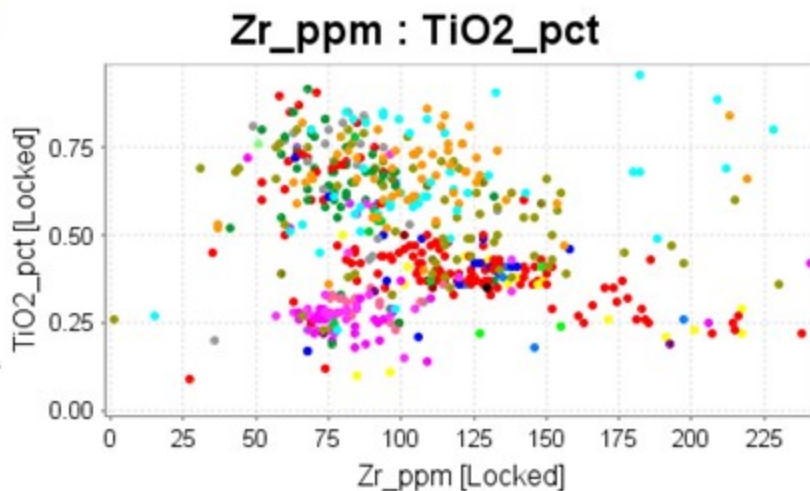
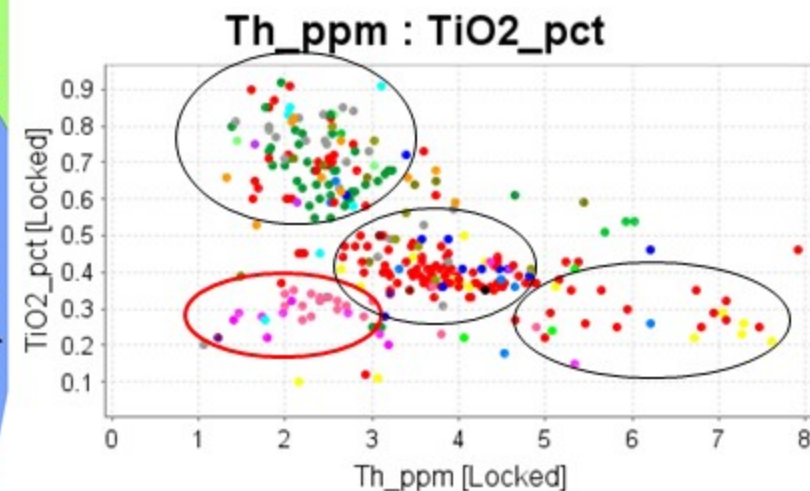
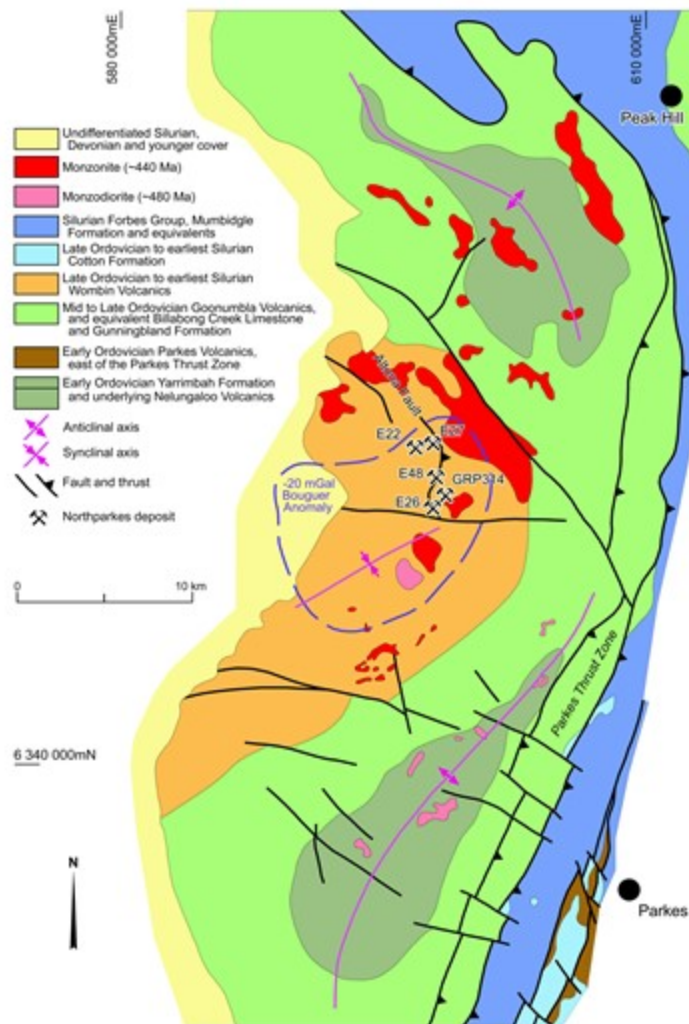
Titanite fractionation removes Nb, Th, Zr, REE's and U from the melt.

Porphyry Cu magmas typically have 0.2 to 0.4% Ti, 2 to 4ppm Th, 2 to 4ppm Nb



Application

In a porphyry district, sample all of the mapped intrusions and classify them geochemically. Identify the prospective suite on your maps. Note there are 4 different magma suites in this data.



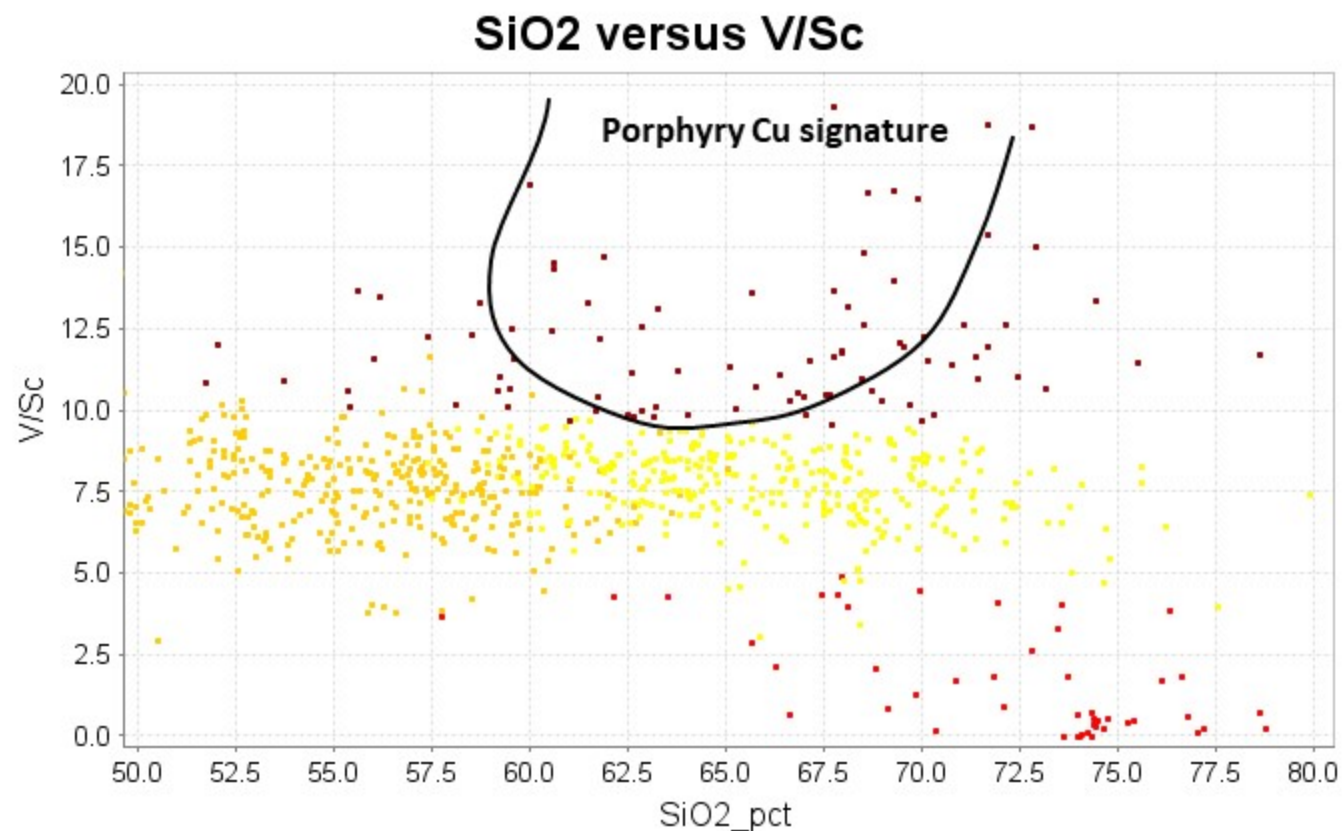
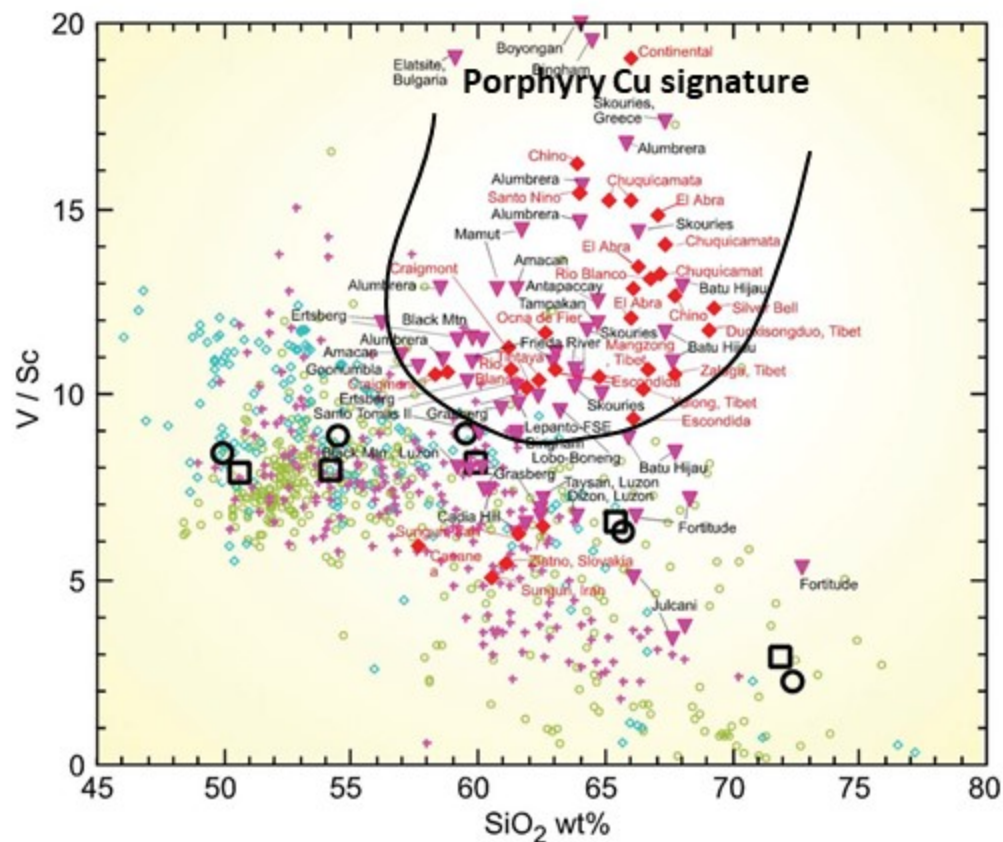
Geological setting of the Goonumbula Volcanic Complex, Macquarie Arc, Central Western New South Wales. After Pacey et al. (2019), Owens et al. (2017), Simpson et al. (2005), Lickfold et al. (2003), Raymond et al. (1995).

Ralph Porter; <http://www.portergeo.com.au/>

Pacey, A., Wilkinson, J.J., Owens, J., Priest, D., Cooke, D.R. and Millar, I.L., 2019. The anatomy of an alkalic porphyry Cu-Au system: geology and alteration at Northparkes Mines, New South Wales, Australia. *Economic Geology*, 114(3), pp.441-472.

Fractionation in Archean Granites

Among the Archean rocks there are signatures of both magnetite fractionation or hornblende fractionation. Increasing V/Sc with increasing SiO₂ is NOT unique to porphyry Cu systems.



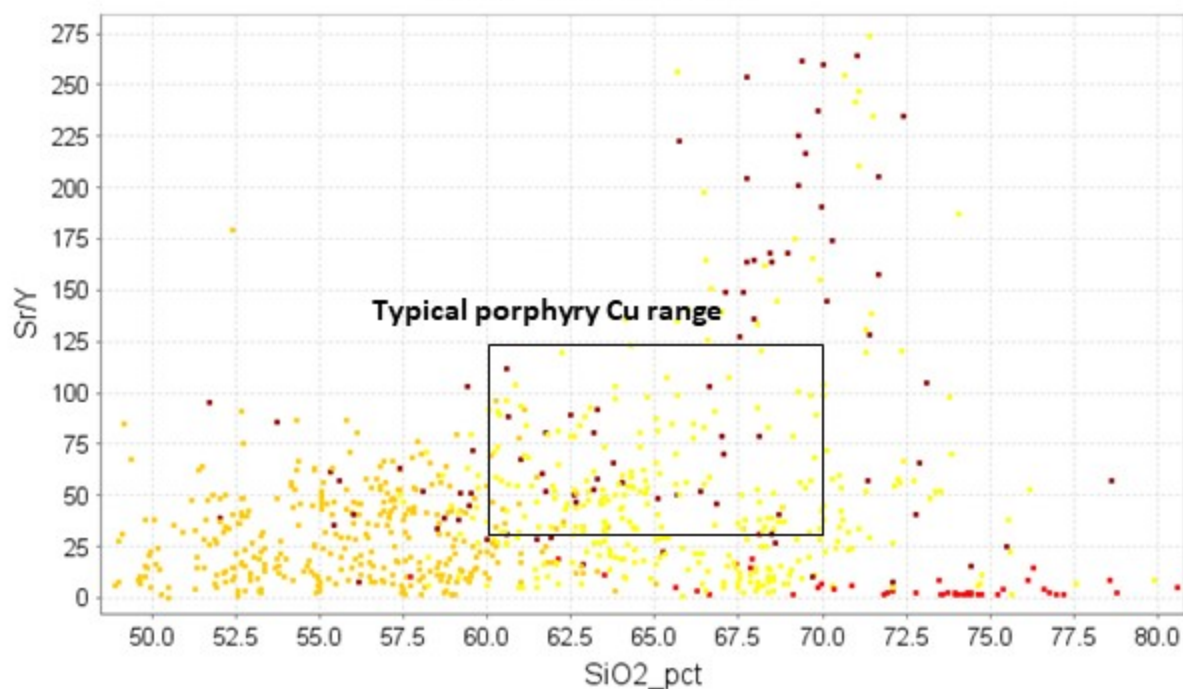
R.R. Loucks (2014) Distinctive composition of copper-ore-forming arc magmas, *Australian Journal of Earth Sciences: An International Geoscience Journal of the Geological Society of Australia*, 61:1, 5-16

Fractionation in Archean Granites

Many Archean granites show the high Sr/Y signature found in porphyry Cu systems; indicative of high pressure melting of a hydrous source region.

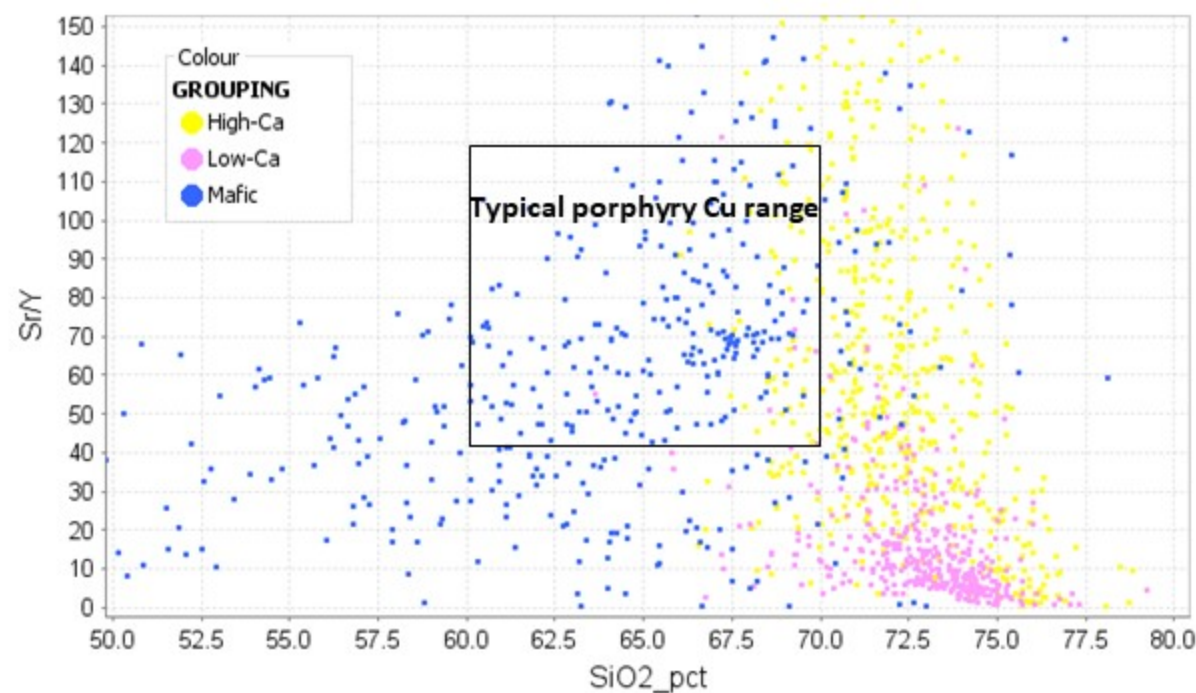
Cassidy and Champion; High Ca granite = high Sr/Y, Low Ca granite = low Sr/Y

SiO₂ versus Sr/Y



Smithies, RH and Lowrey, JR 2020, Eastern Goldfields greenstone geochemical barcoding project: notes to accompany 2020 data release: Geological Survey of Western Australia, Record 2020/6, 4p

SiO₂ versus Sr/Y



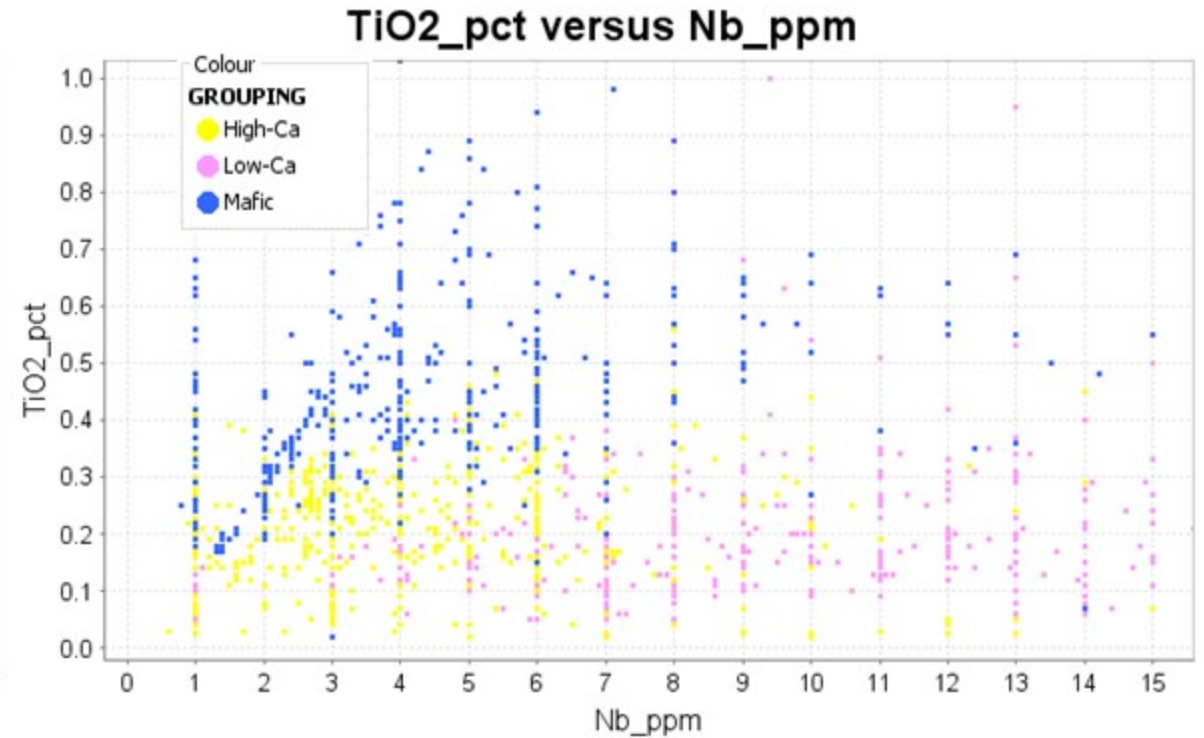
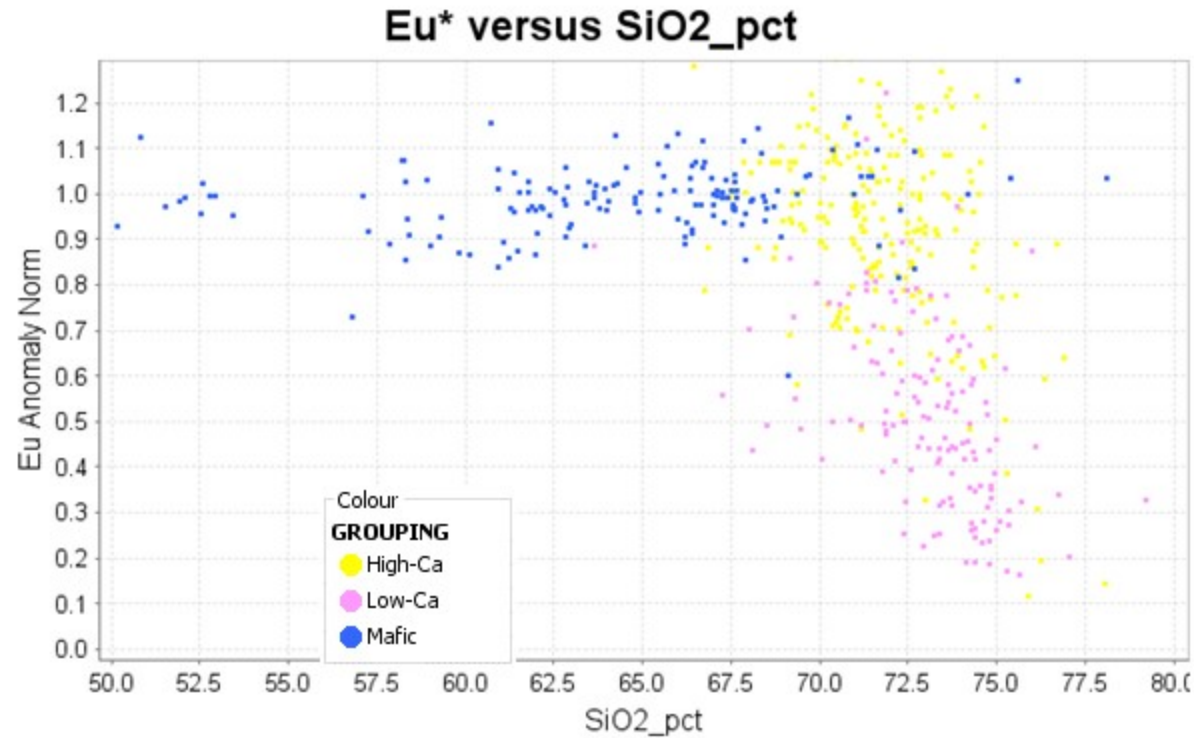
AMIRA P482. Kevin Cassidy and Dave Champion, Yilgarn Granites

Fractionation in Archean Granites

High Sr/Y granite has NOT fractionated plagioclase and has low abundances of Nb, Th, Zr, REE's, U etc.

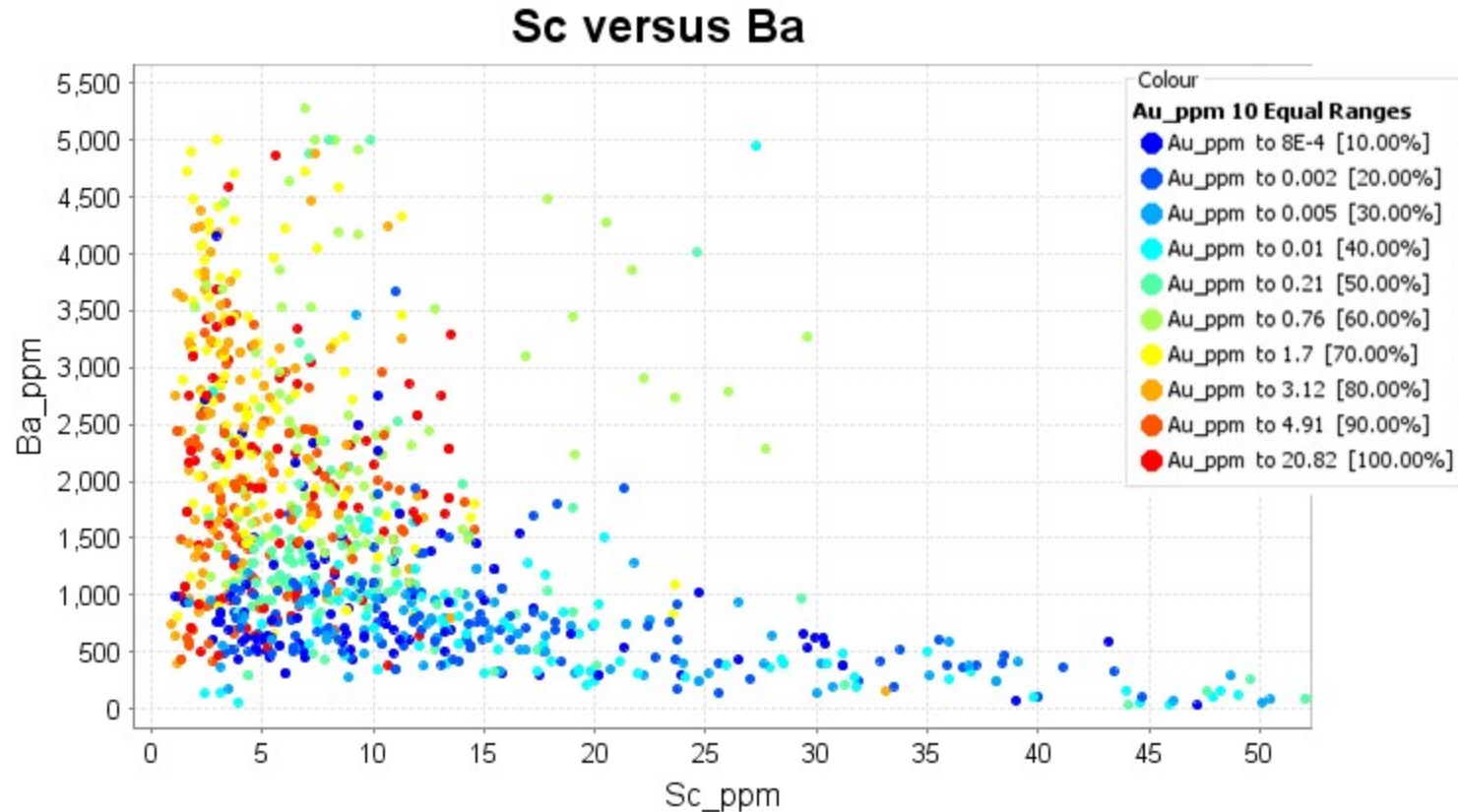
Low Sr/Y granite shows incompatible behaviour of HFSE's

Titanite fractionation; same pattern as in porphyry Cu magmas.



Fractionation in Archean Granites

Very distinct spatial association of hornblende diorites with gold deposits, where those diorites have fractionated hornblende, titanite, but NOT plagioclase, AND where the magmas are sourced from a metasomatized mantle. Signature of this is $Ba > 1000 \text{ ppm}$.



Eastern Goldfields; >20 diorite intrusions with 1moz to 40moz Au at grades of 0.2 to 0.5ppm.

- Very sulfur-rich!!
- Intrinsically Au-rich
- Distinct trace element signatures
- Partial melts of metasomatised mantle,



Beta Porphyry, Kambalda. Pyrite +/- anhydrite.
Mean gold grade 0.37ppm Au



Santa Anna Trondhjemite, Kambalda..
Mean gold grade 0.5ppm Au

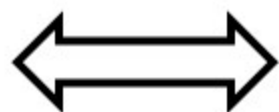
Fractionation in Yilarn Granites

Very distinct spatial association of hornblende diorites with gold deposits.

Low-Ca granites

Low pressure mid-crustal melting/fractionation

- High K melt
- Low Ca-Sr
- Decreasing Sr/Y with increasing SiO₂
- Eu* anomaly
- Decreasing V/Sc
- Incompatible behaviour of REE's, Nb, Th, Zr, U



High-Ca granites

High pressure melting/fractionation of oxidized hydrated mafic

- High Ca-Sr melt
- Increasing Sr/Y with increasing SiO₂
- No Eu* anomaly
- Increasing V/Sc
- Compatible behaviour of REE's, Nb, Th, Zr, U



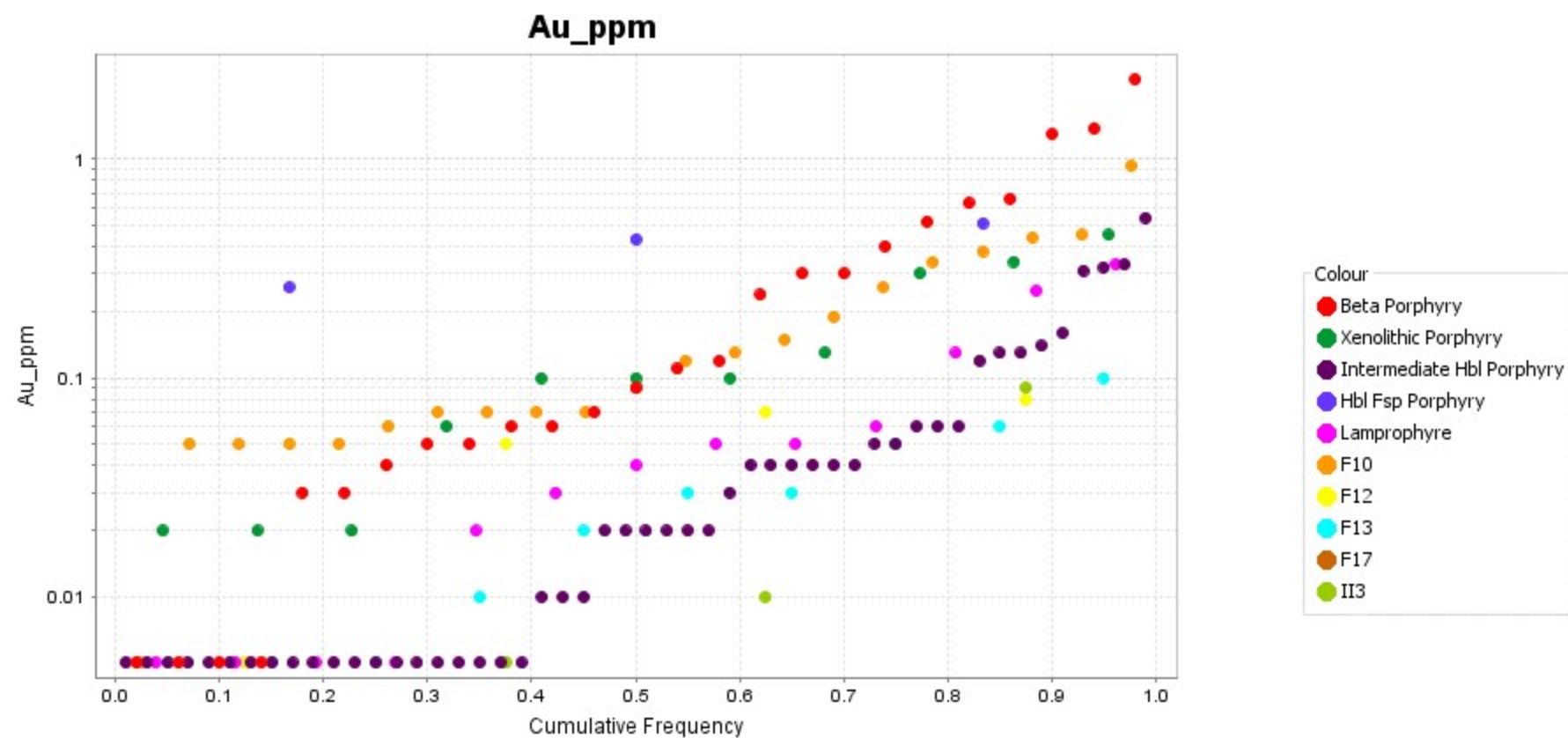
Alkalic (Syenites)

Low degree partial melting of metasomatized source

- High K-Na melt
- High Sr and Ba, but also high Y
- No Eu* anomaly
- Very high V/Sc
- Strongly enriched in REE's, Nb, Th, Zr, U
- Zr/Hf commonly up to 50

Yilgarn Granites

diorite intrusions with 1moz to 40moz Au at grades of 0.2 to 0.5ppm.



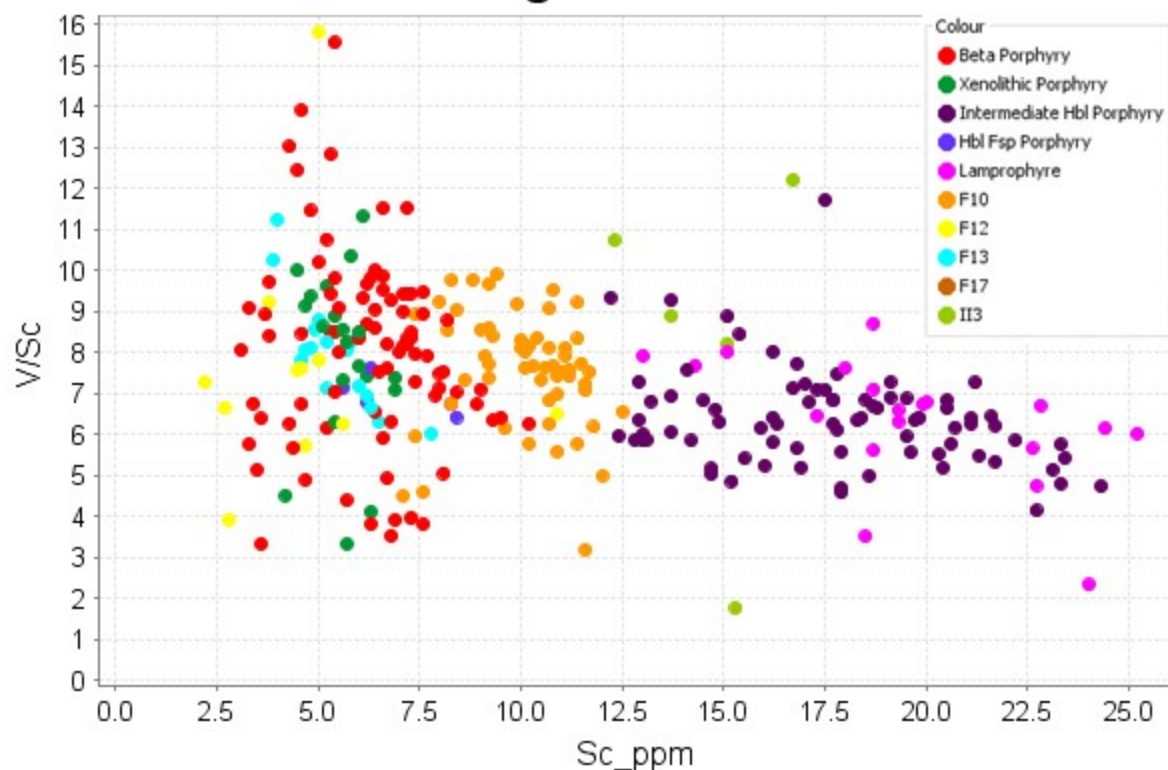
Compare Yilgarn Granites with porphyry Cu magmas

Low Ca-granites have declining V/Sc; magnetite fractionation

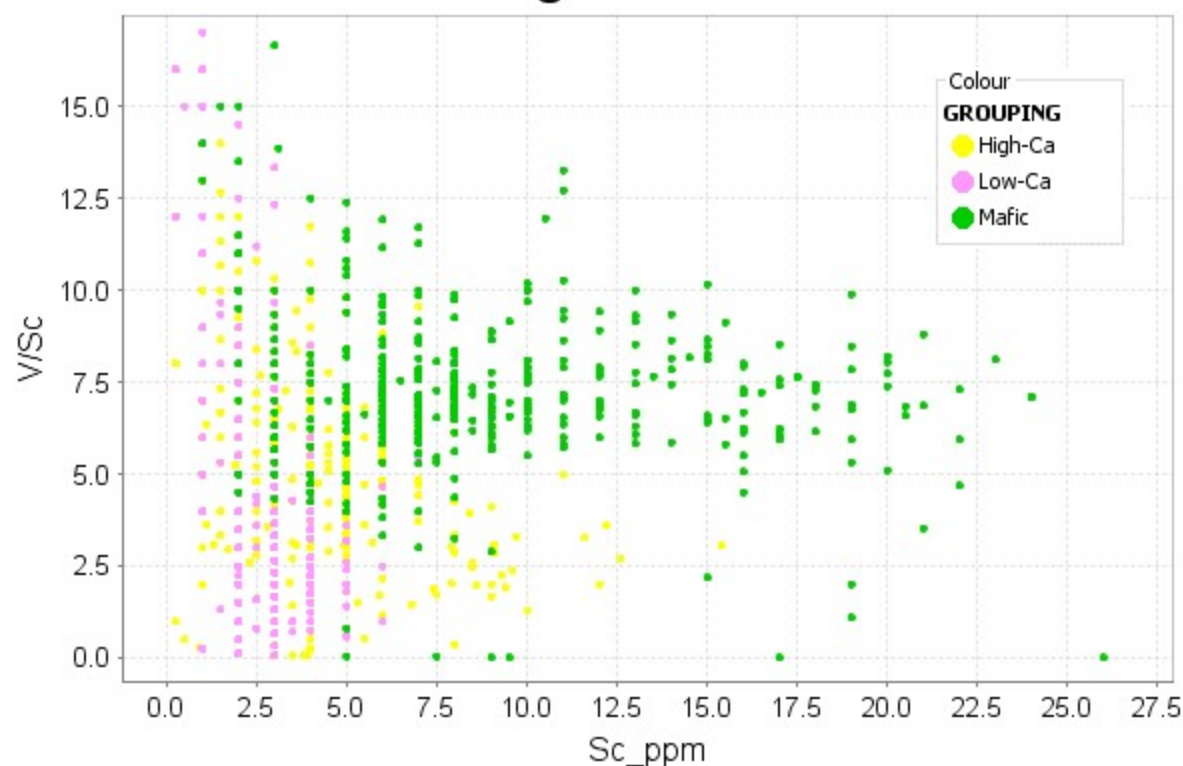
Some high Ca-granites have increasing V/Sc; hornblende fractionation

Beta porphyry (red) is a mineralized intrusion at St Ives.

Hornblende-Magnetite Fractionation



Hornblende-Magnetite Fractionation



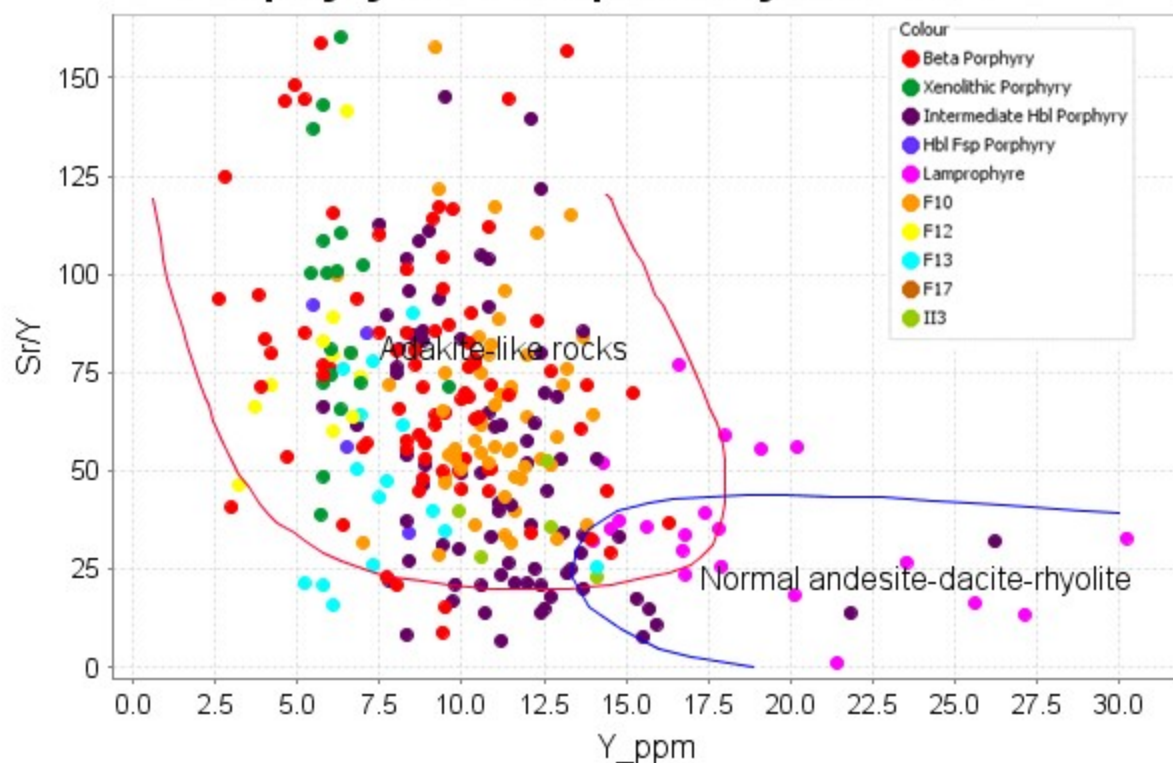
Compare Yilgarn Granites with porphyry Cu magmas

Low Ca-granites have low Sr/Y (Sr declines with plagioclase fractionation)

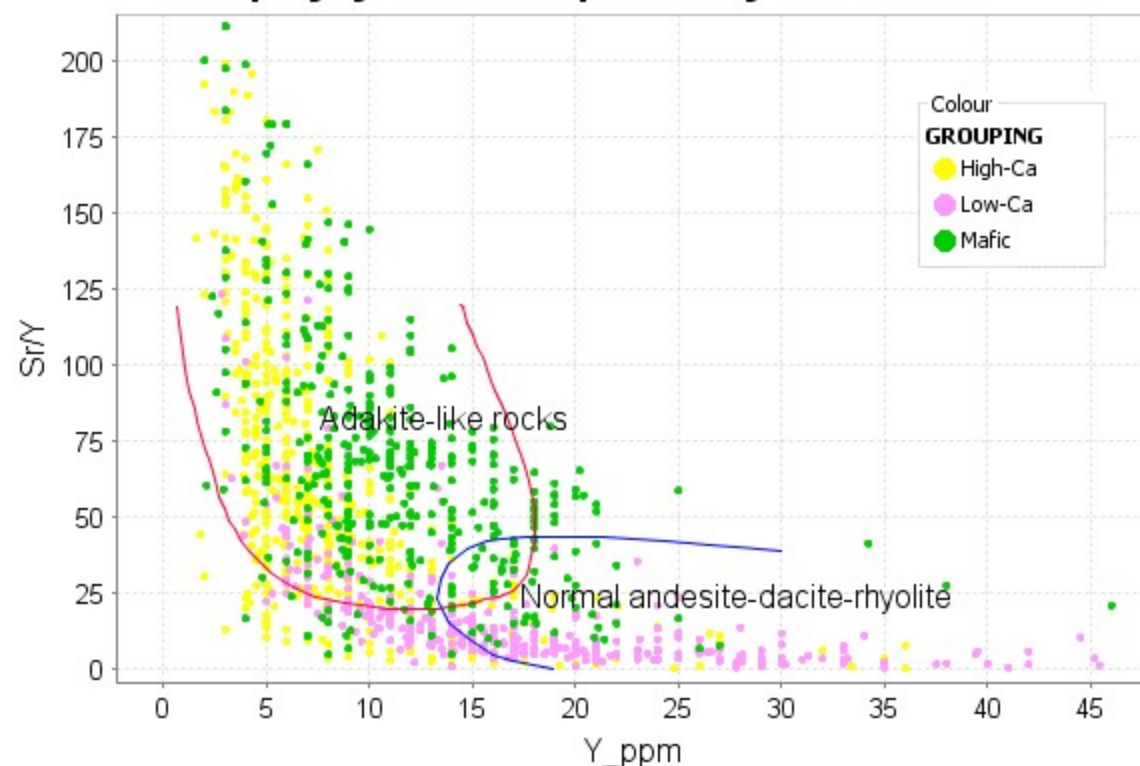
High Ca-granites have increasing Sr/Y (titanite fractionation before plagioclase crystallization)

Beta porphyry (red) is a mineralized intrusion at St Ives.

Porphyry Cu Prospectivity Sr/Y vs Y



Porphyry Cu Prospectivity Sr/Y vs Y

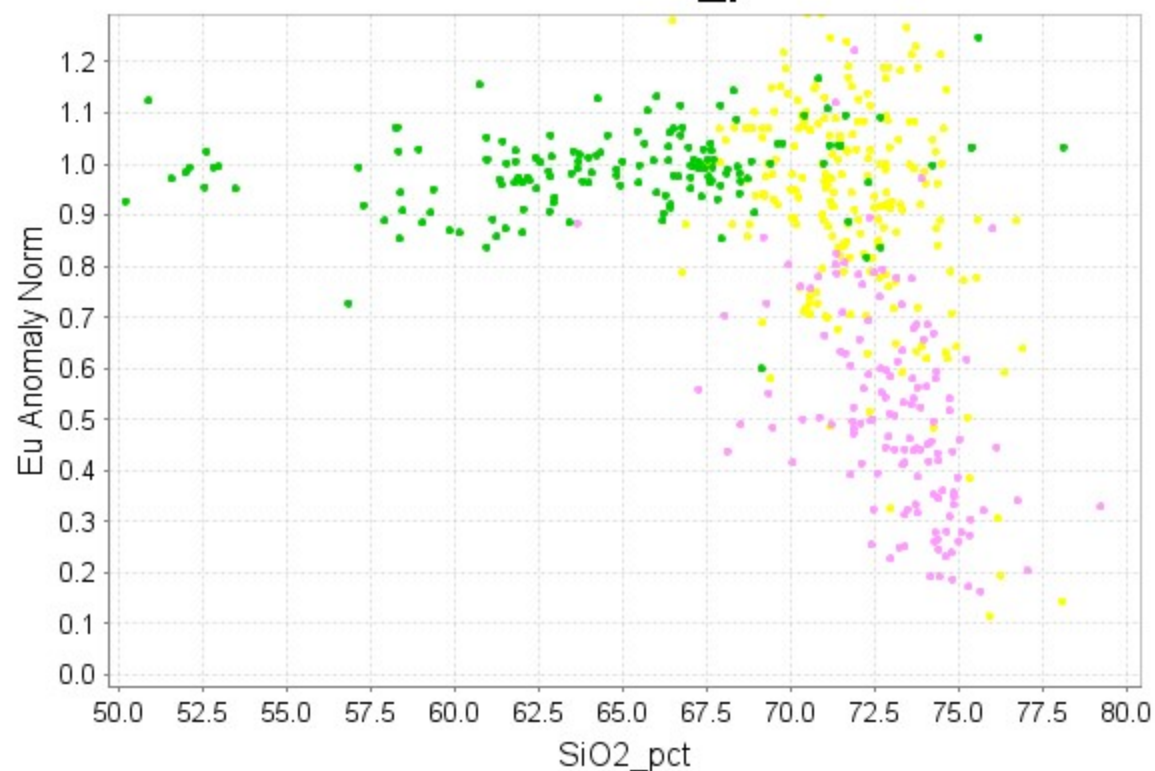


Compare Yilgarn Granites with porphyry Cu magmas

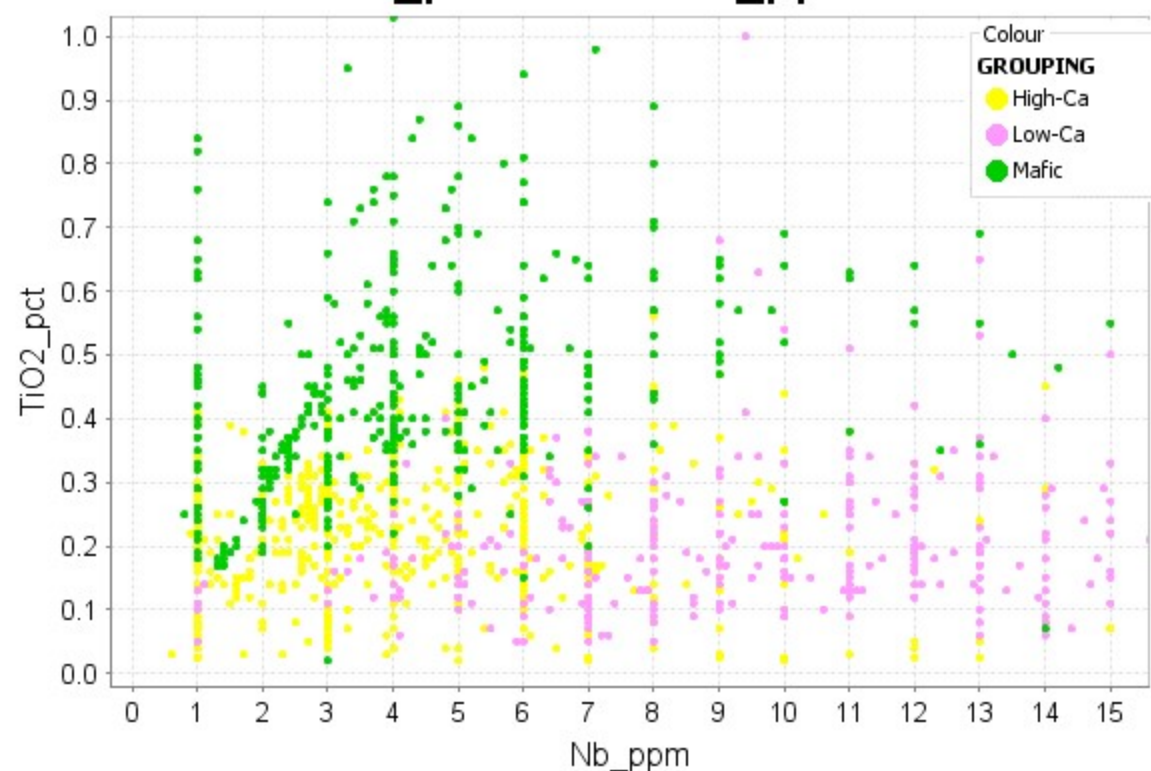
Low Ca-granites have negative Eu anomaly AND incompatible HFSE's

High Ca-granites have NO Eu anomaly AND compatible HFSE's (titanite fractionation)

Eu* versus SiO₂_pct



TiO₂_pct versus Nb_ppm

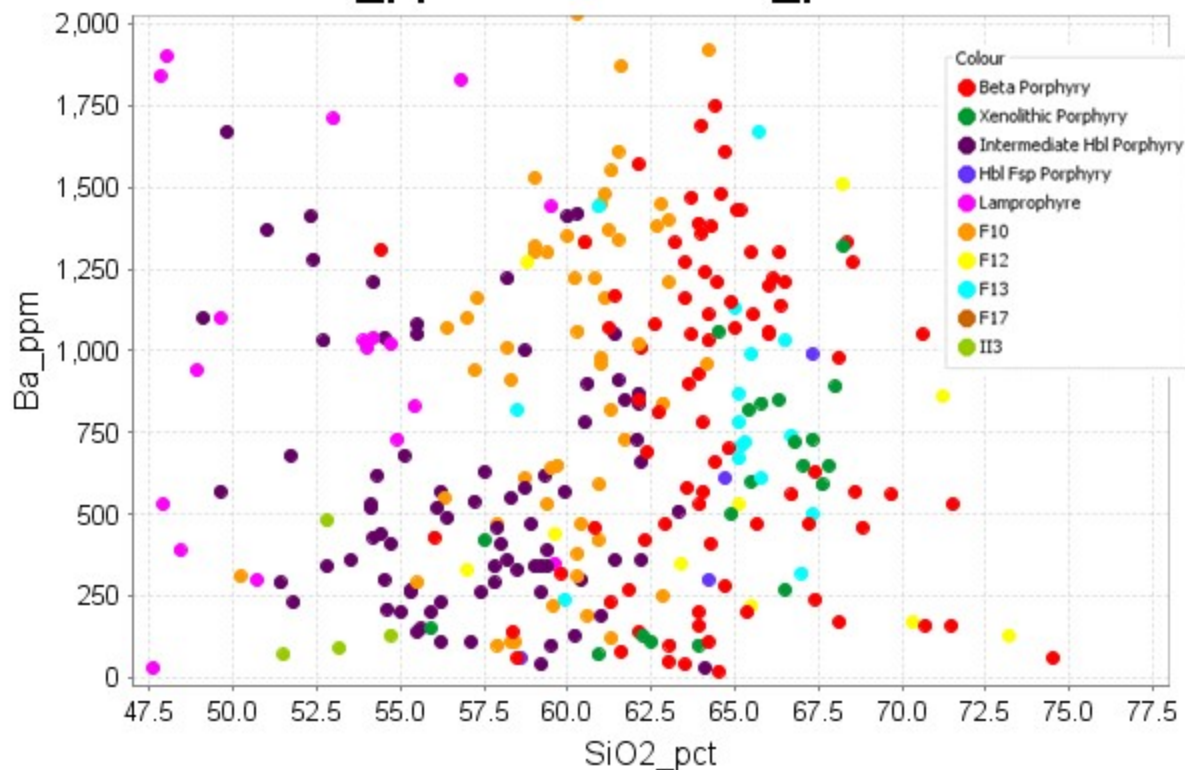


Compare Yilgarn Granites with porphyry Cu magmas

Low Ca-granites usually have high Ba (high K, high Ba)

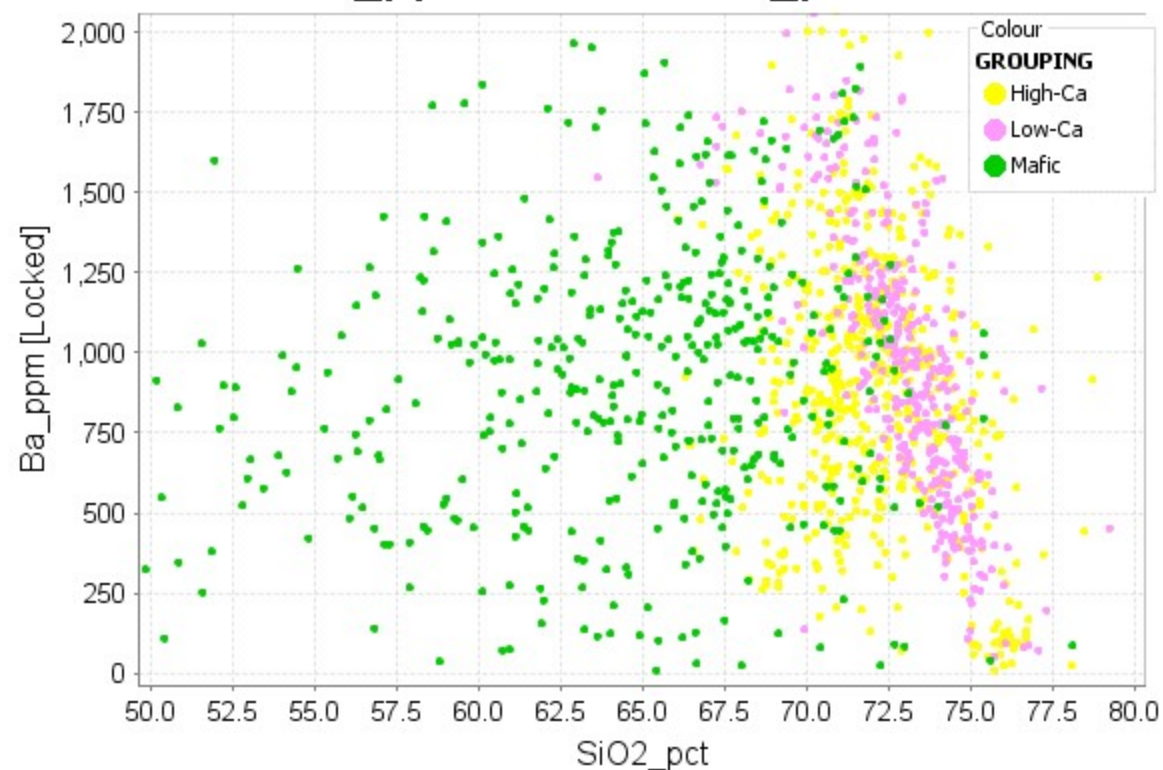
Mineralized High Ca-granites have high Ba, indicative of a metasomatized mantle source!

Ba_ppm versus SiO₂_pct



Dave Champion, Geoscience Australia, St Ives porphyry study, 2008

Ba_ppm versus SiO₂_pct



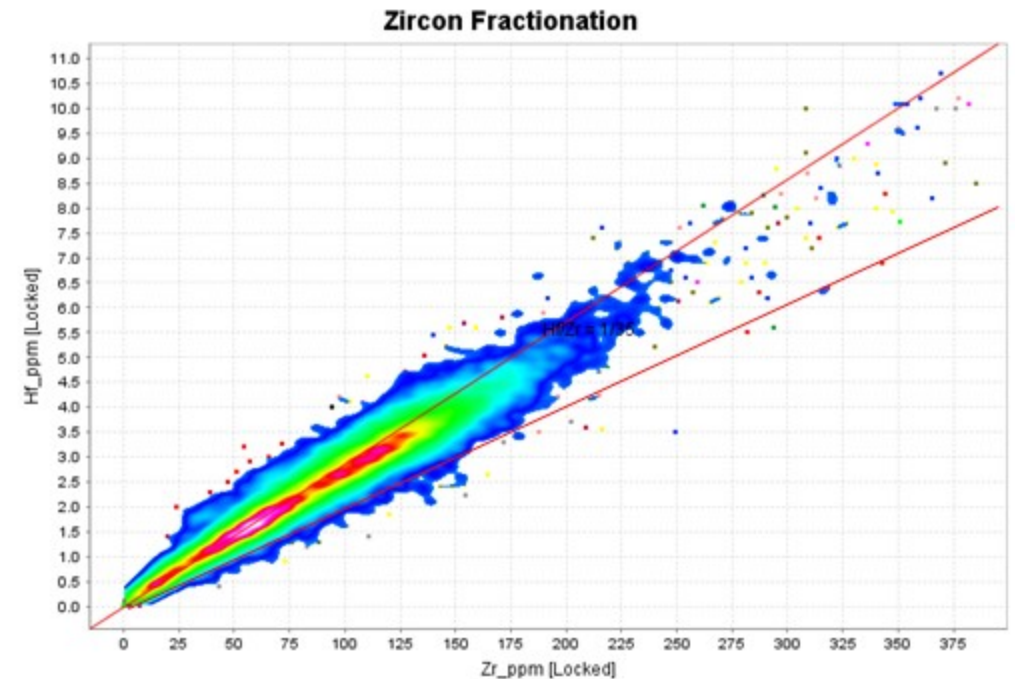
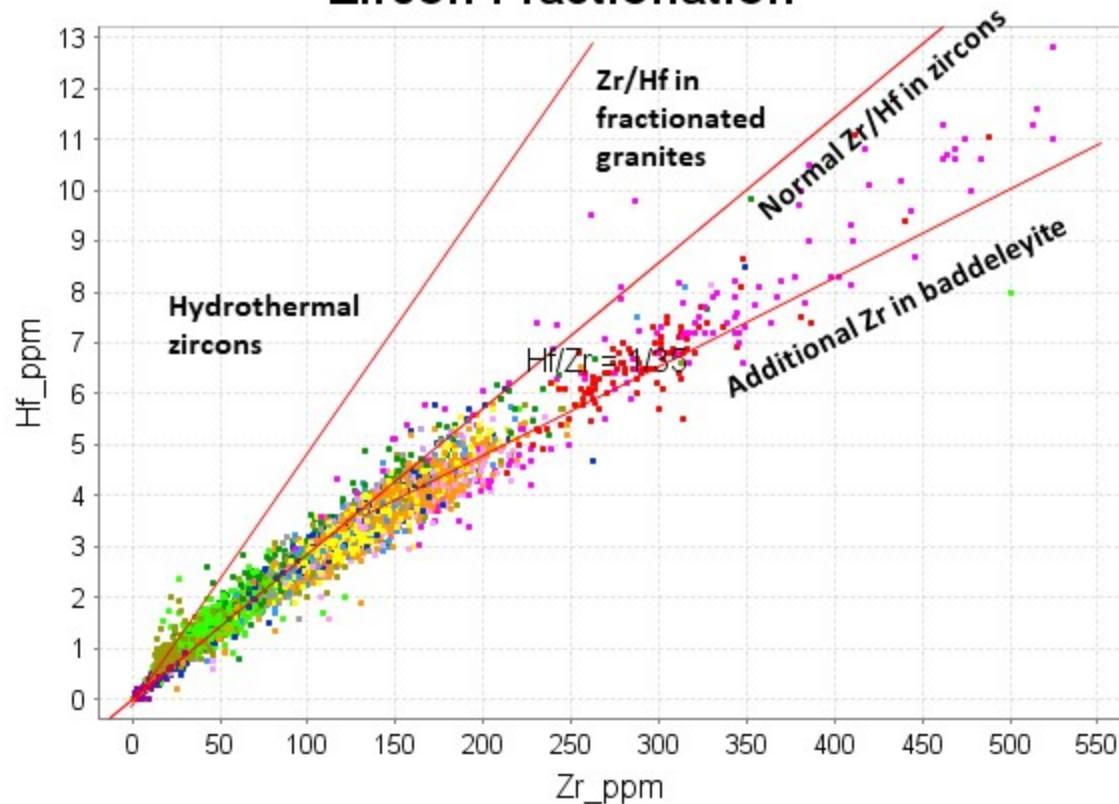
AMIRA P482. Kevin Cassidy and Dave Champion, Yilgarn Granites

Alkalic Magmas

There is a relatively rare compositional group of Archean gold-related intrusions classified as syenites. These have;

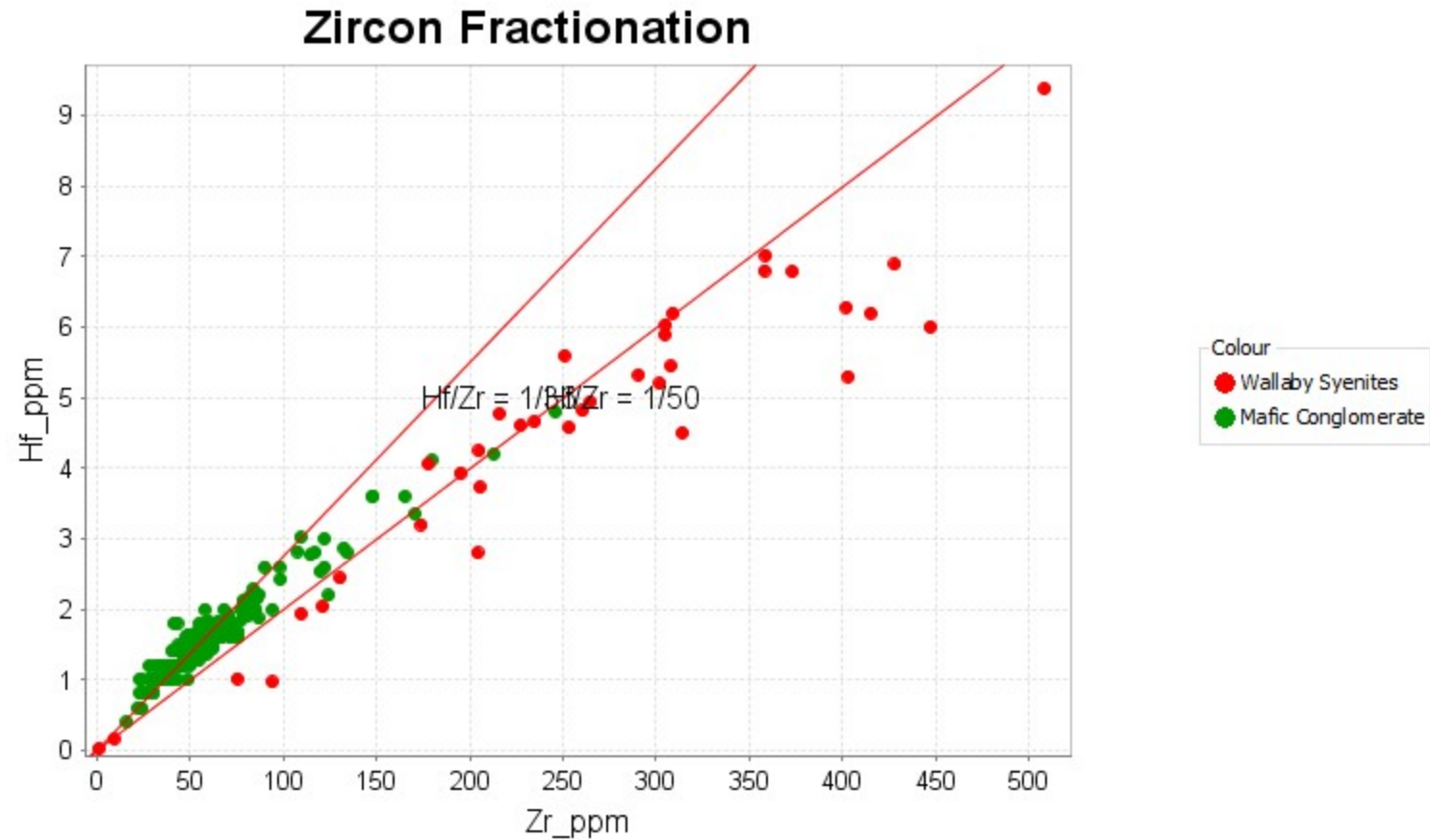
- increasing V/Sc with increasing SiO₂,
- strong enrichment in P, LREE's, Th and Zr
- Ba>1000ppm, enrichment in Sr but not depletion in Y.
- Unique signature of zircon + baddeleyite, increases Zr/Hf ratio from background of 36 up to 50.

Zircon Fractionation



Alkalic Magmas

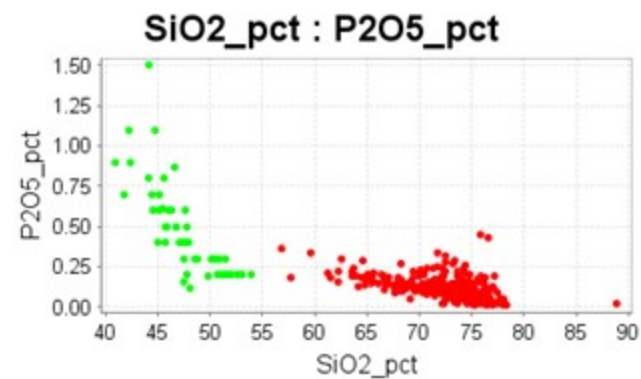
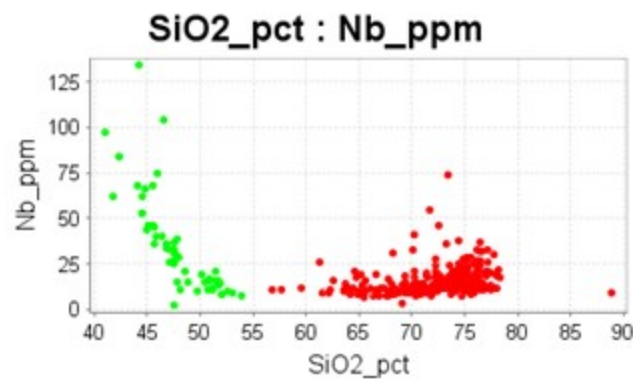
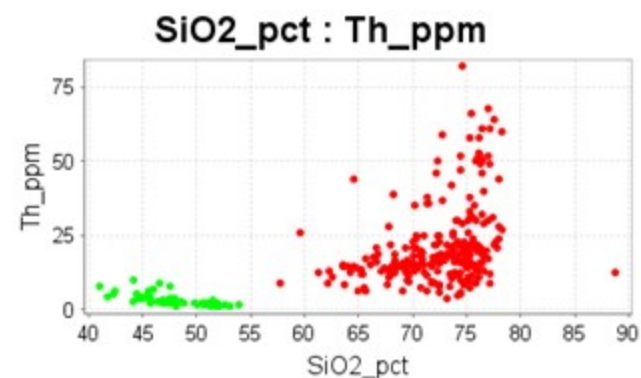
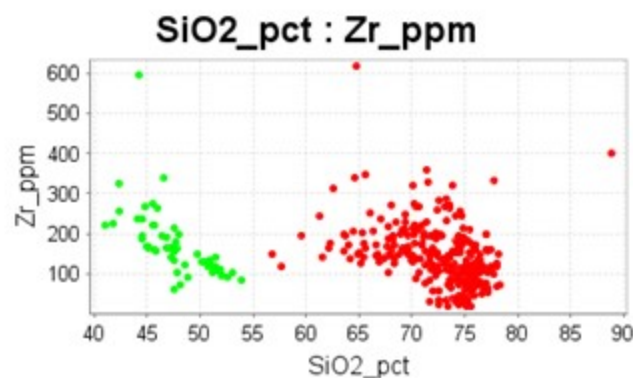
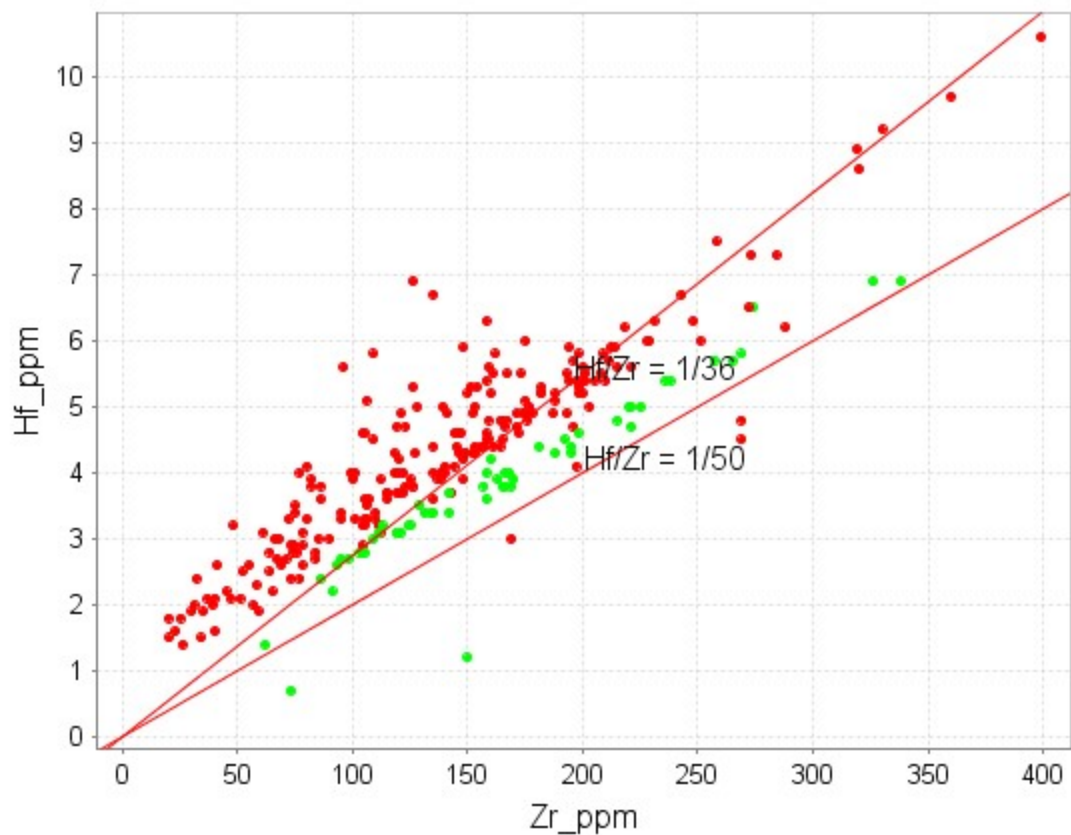
Most notable examples of Alkalic systems are Wallaby (Laverton) and Porgera



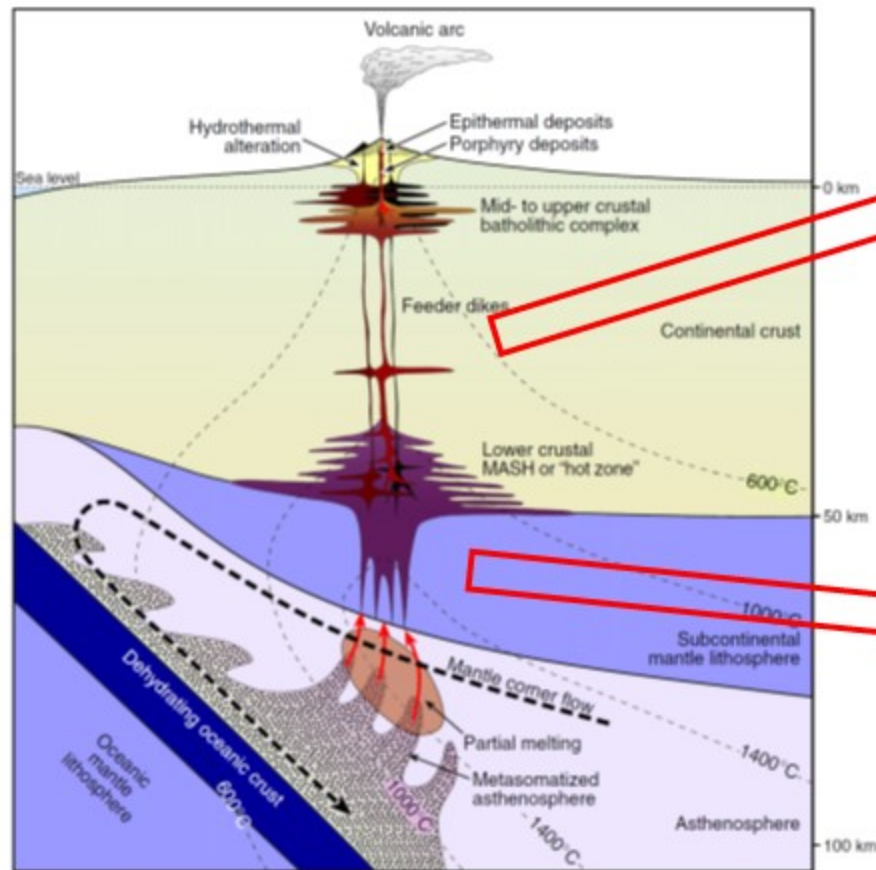
Alkalic Magmas

Tasmanian Devonian Granites (red) and Tertiary Alkali Basalts (green)

Zircon Fractionation



Summary



Low pressure mid-crustal melting/fractionation

- High K melt
- Low Ca-Sr
- Decreasing Sr/Y with increasing SiO₂
- Eu* anomaly
- Decreasing V/Sc
- Incompatible behaviour of REE's, Nb, Th, Zr, U

High pressure melting/fractionation of oxidized hydrated mafic

- High Ca-Sr melt
- Increasing Sr/Y with increasing SiO₂
- No Eu* anomaly
- Increasing V/Sc
- Compatible behaviour of REE's, Nb, Th, Zr, U

Conclusions

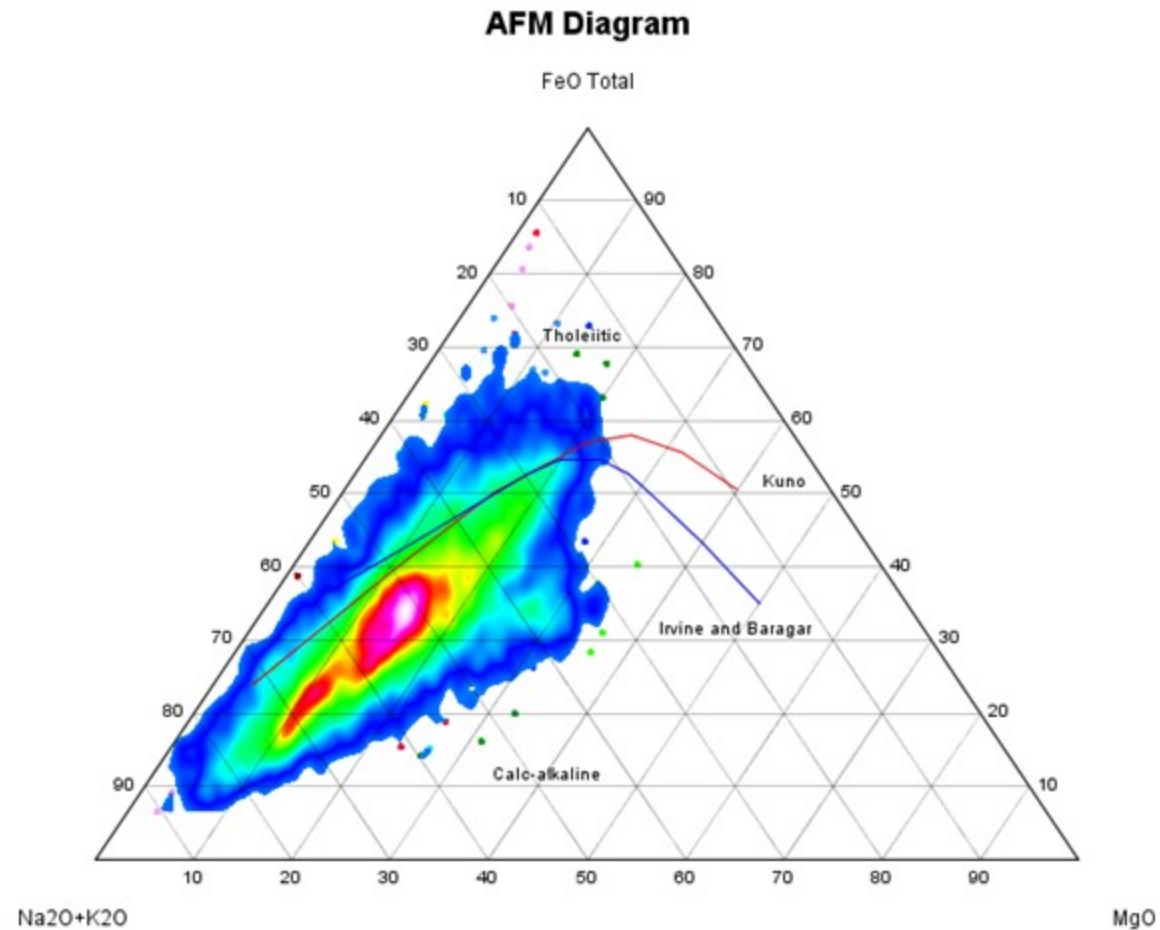
- Exsolution of volatiles from fractionating magmas is a critical process in many hydrothermal mineral systems.
- Trace element pairs with a common mineral host are sensitive indicators of fractionation.
- The order and degree of mineral fractionation is determined by source composition, pressure, temperature, water content, oxidation state, sulfide saturation etc. All of these impact on the potential to create a mineral deposit.
- Exploration and Mine geologists can use routine 4 acid digest ICP-MS data to create the plots presented here to understand the nature of igneous rocks in their drill holes.

References

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Conclusions

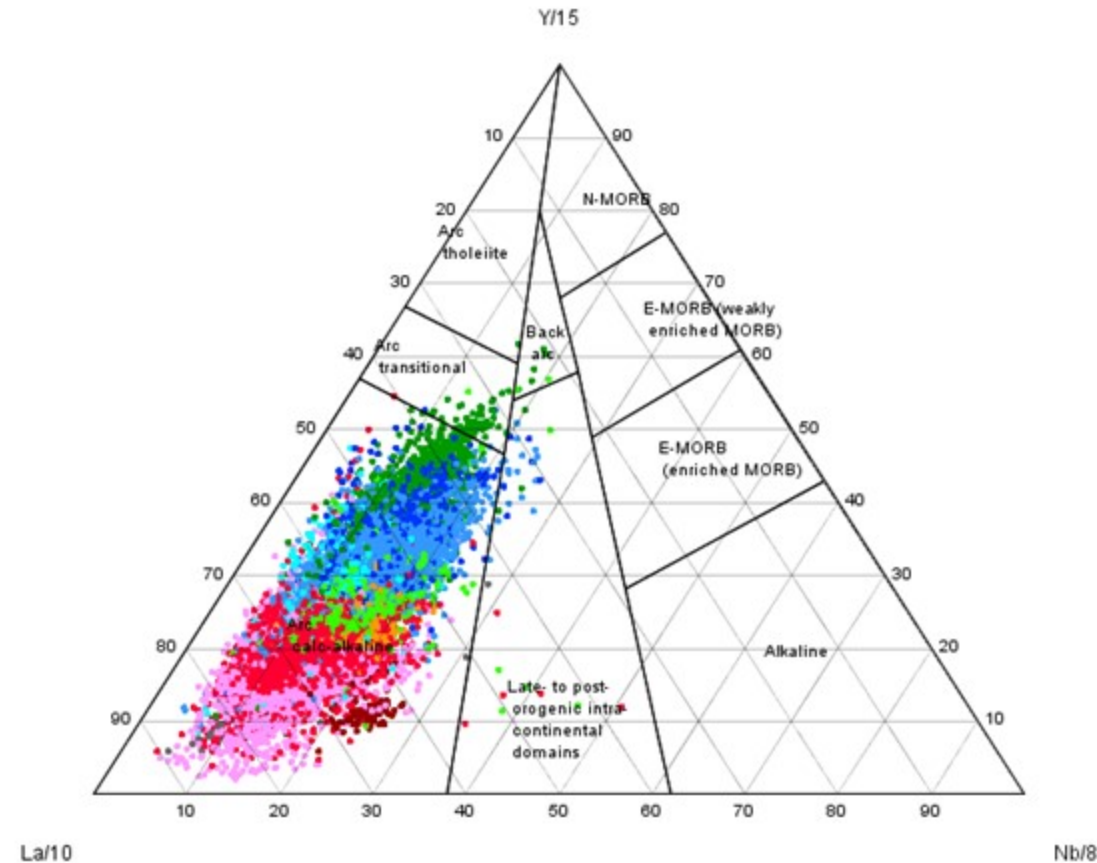
- Don't use major element diagrams to classify compositions of altered rocks!



Conclusions

- Use 1980's style trace element classifications with caution! Many of these are ambiguous. Many are designed to compare igneous series at comparable SiO₂ contents

Tectonic Classification of Mafic Igneous Rocks (Cabanis and Lecolle 1989)



Conclusions

- Spider diagrams are way too confusing. Don't use statistical methods. Use the actual assay numbers and interpret these in terms of the host mineral compositions.

